

## Origin and Biogeographic History of *Scapharca broughtonii* (Schrenck, 1867) (Bivalvia: Arcidae) and Its Related Species

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**Abstract:** *Scapharca broughtonii* (Schrenck, 1867) and *S. aff. broughtonii* have been recovered for the first time from the upper Pliocene (Piacenzian) deposits on the Japan Sea side of Honshu. This is the oldest record of *S. broughtonii* and its related species. *S. satowi* is known from the same locality as *S. broughtonii* in the upper Pliocene Tentokuji Formation in Akita Prefecture, Japan Sea side of northern Honshu. These fossils suggest that the genetically similar ark shells speciated near the northern limit of the warm-water current in the semi-enclosed Japan Sea by the late Pliocene. In the middle Pleistocene (Chibanian), *S. broughtonii* first adapted to cold-temperate water, being known from the Dateyama Formation on the Japan Sea side of central Hokkaido and associated with many cold-water and a few temperate-water species.

**Keywords:** late Pliocene, Tentokuji Formation, warm-water current, cold-water adaptation

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### Introduction

The red ark shell, *Scapharca broughtonii* (Schrenck, 1867), lives in sandy and muddy bottoms from subtidal to 60 m in depth and is widely distributed from the southern Primorie and southern Sakhalin areas of Russia through Hokkaido to Kyushu, Korea, the Bohai Sea, Yellow Sea and East China Sea (Akamatsu, 1992; Higo *et al.*, 1999; Scarlato, 1981; Qi *et al.*, 1989; Li, 2004; Matsukuma & Okutani, 2000, 2017; Xu & Zang, 2008; Lutaenko & Noseworthy, 2012; Zhang *et al.*, 2016). This is one of the more important bivalves in fisheries in Japan and adjacent countries (*e.g.*, Tamura, 1960). Its type locality is Hakodate in southern Hokkaido, northern Japan (Schrenck, 1867). Most species of *Scapharca* live in mild-temperate areas (Nishimura, 1981) and southwards (*e.g.*, Higo *et al.*, 1999) but *S. broughtonii* is unusual in also inhabiting cool-temperate and subarctic areas (Nishimura, 1981) such as the Bohai Sea and the southern Primorie Region of Russia. However, it is uncertain when and where the species first appeared and adapted to cold-water areas.

Yokogawa (1997) noted that the Chinese population from Dalian identified as *S. broughtonii* has more numerous radial ribs, ranging from 43 to 50 (average 45.83), whereas the number in the Japanese population ranges from 39 to 44 (average 41.25) including the type specimens (42 to 43). He also recognized a genetic difference between two forms. Based on examination of the mt DNA COI, Cho *et al.* (2007) confirmed that the Korean and Russian red ark shells are genetically different from the Chinese (Shandong) population. Tanaka & Aranishi (2016) separated two lineages, the Korean and Japanese Lineage A and Chinese (Dalian) Lineage B. From these results, *S. broughtonii* can be separated into two species or subspecies as claimed by Yokogawa (1997).

Fossil records of *S. broughtonii* are common from the Pleistocene of the Pacific side of central

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Honshu (Baba, 1990). Ozawa *et al.* (1998) described this species from the lower Pleistocene (Gelasian) Dainichi Formation (2–1.9 Ma) and lower Pleistocene (Calabrian) Aburayama Formation in Shizuoka Prefecture. Jimbo *et al.* (2017) illustrated *S. broughtonii* from the lower Pleistocene (Gelasian) Imaizumi Sandstone and Conglomerate Member of the Nojima Formation in Kanagawa Prefecture whose age has been established as 2.5–1.97 Ma based on magnetostratigraphy. These early Pleistocene occurrences were associated with the warm-water Kakegawa fauna (Otuka, 1939b). On the Japan Sea side of Honshu, on the other hand, fossils of this species are not so common. The oldest specimen on the Japan Sea side was reported from the lowermost Pleistocene (Gelasian) part (dated at around 2.4 Ma by Yanagisawa & Amano, 2003) of the Nadachi Formation at Osuga in Joetsu City, Niigata Prefecture by Amano *et al.* (1988). From this locality, the Omma-Manganji fauna was collected (Otuka, 1939a), consisting of some cold-water and extinct endemic species with fewer warm-water species.

*Scapharca broughtonii* and *S. aff. broughtonii* were recovered for the first time from upper Pliocene deposits on the Japan Sea side of Honshu that are older than the Nadachi Formation and the Nojima Formation. In this paper, we describe these late Pliocene specimens and discuss the biogeographic history of this species by considering climatic change and the geohistory of the Japan Sea.

## Material and Methods

*Scapharca broughtonii* has been newly recovered from the following three localities (Locs. 1–3; Fig. 1); three specimens from the upper Pliocene (Piacenzian) Tentokuji Formation at a large cliff along Yodo River, Awasegai in Daisen City, Akita Prefecture, northern Honshu (Loc. 1 = Loc. 21 of Amano *et al.*, 2000; 39°37′14″N, 140°18′41″E), one specimen from the lowermost Pleistocene (Gelasian) Nadachi Formation (uppermost part) at the river bed of the Nadachi River near Osuga Bridge in Joetsu City, Niigata Prefecture, central Honshu (Loc. 2; 37°8′46″N, 138°6′20″E) and two specimens from the uppermost lower Pleistocene (Calabrian) part of the Omma Formation (upper part of the formation *sensu* Kitamura & Kondo, 1990) at 450 m upstream from Okuwa (formally called Omma) Bridge in Kanazawa City, Ishikawa Prefecture (Loc. 3; 36°32′6″N, 136°40′50″E). The molluscan fossils from the last locality correspond with the Younger Ommaian fauna (Ogasawara, 1981).

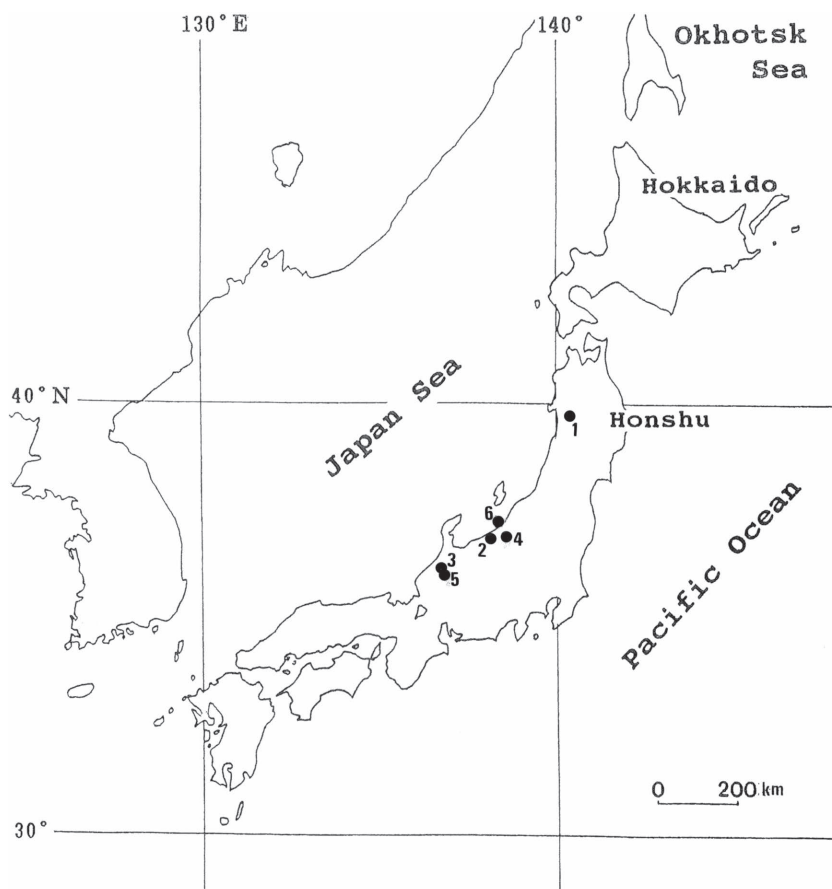
Two specimens of *S. aff. broughtonii* were collected from the upper Pliocene (Piacenzian) Higashigawa Formation exposed along the Shibumi River between Tazawa and Inubushi in Tokamachi City, Niigata Prefecture, central Honshu (Loc. 4; 37°8′22″N, 138°38′23″E).

For comparison we examined 56 specimens of *Scapharca ommaensis* (Otuka, 1936) from the lower Pleistocene (Calabrian) part of the Omma Formation on the bank of the Saikawa River, 1 km upstream of Okuwa Bridge in Kanazawa City, Ishikawa Prefecture (Loc. 5; 36°31′51″N, 136°40′59″E) and 31 Recent specimens of *Scapharca satowi* Dunker, 1882 collected from the Kakizaki coast in Joetsu City, Niigata Prefecture (Loc. 6; 37°16′55″N, 138°23′21″E). The terminology follows Reinhart (1943) and Noda (1966). All specimens treated herein are stored in the Department of Geology and Paleontology (NMNS PM) of the National Museum of Nature and Science, Tsukuba.

## Description of the Pliocene species

Family **Arcidae** Lamarck, 1809

Subfamily **Anadarinae** Reinhart, 1935



**Fig. 1.** Locality map of fossils and Recent specimens. See the text on the information of localities.

Genus *Scapharca* Gray, 1847

**Type species:** *Arca inaequalvis* Bruguière, 1837 (original designation)

***Scapharca broughtonii* (Schrenck, 1867)**

(Fig. 2)

*Arca inflata* Reeve, 1844: *Arca* sp. 30; Tokunaga, 1906: 57, pl. 3, fig. 19; Yokoyama, 1920: 167, pl. 17, fig. 9; Kinoshita & Isahaya, 1934: 12, pl. 9, fig. 65.

*Arca reeveana* Nyst, 1848: 60.

*Arca broughtonii* Schrenck, 1867: 578, pl. 24, figs 1–3.

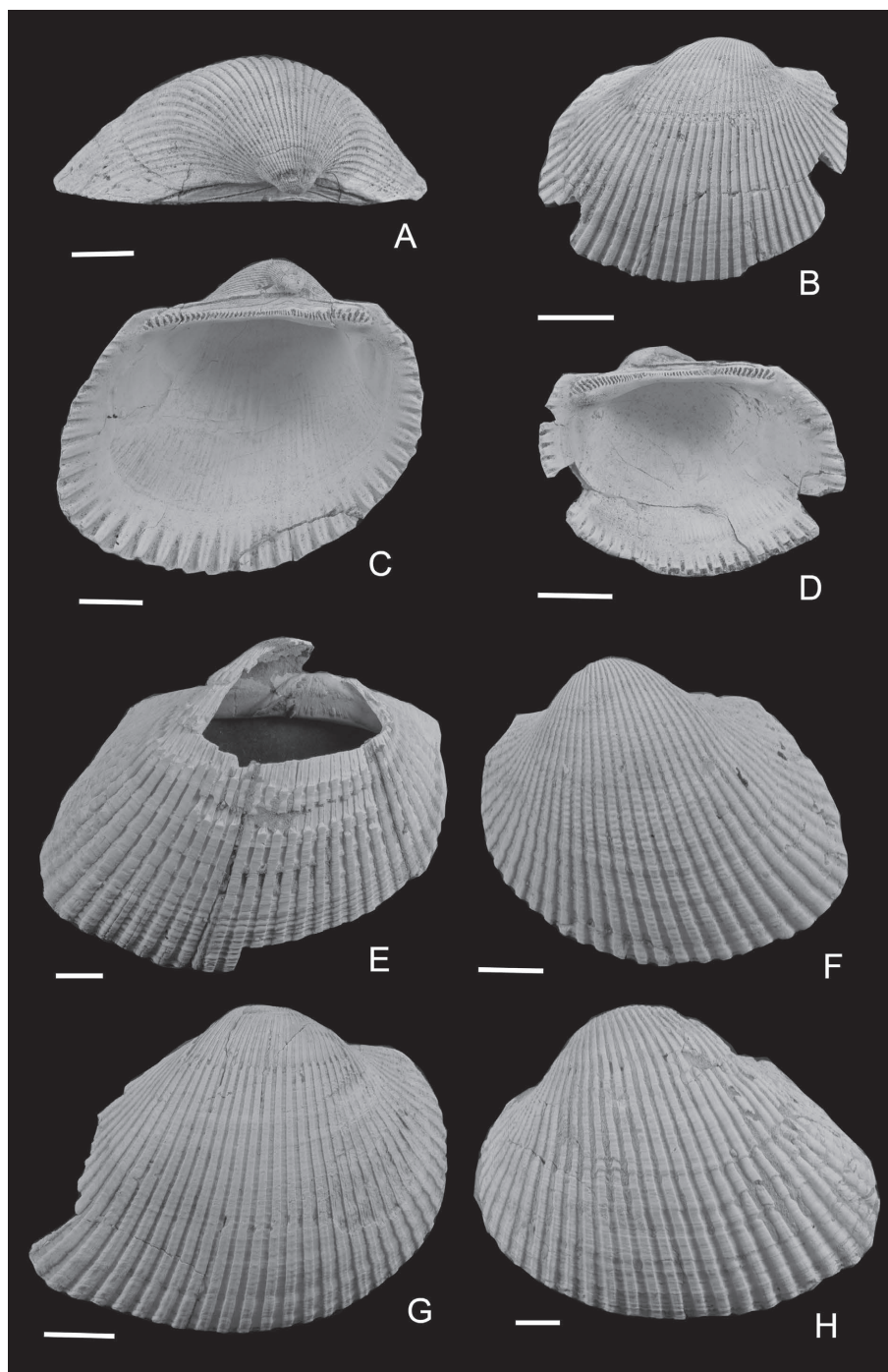
*Arca (Anomalocardia) inflata* — Kobelt, 1891: 30, pl. 10, figs 1, 2; Yokoyama, 1922: 187, pl. 15, fig. 9.

*Arca tenuis* Tokunaga, 1906: 58, pl. 4, fig. 1.

*Arca (Scapharca) inflata* — Yamakawa, 1911: 17, pl. 4, fig. 13, pl. 5, figs 1–4, pl. 6, figs 1–6, pl. 7, figs 1–5.

*Anadara (Scapharca) inflata* — Kuroda, 1930: 32.

*Anadara (Scapharca) broughtonii* — Fujie, 1958: 666, pl. 28, fig. 26; Kira, 1959: 112, pl. 43, fig. 14; Kira, 1962: 124, pl. 4, fig. 14; Noda, 1966: 107; Kobayashi, 1976: pl. 1, figs 1–4; Matsuura,



**Fig. 2.** *Scapharca broughtonii* from the upper Pliocene to lower Pleistocene deposits in the Japan Sea side of Honshu. **A–D, F, G.** Fossils from the upper Pliocene (Piacenzian) Tentokuji Formation at Loc. 1; **A, C, F,** NMNS PM 65051-1, left valve; **B, D,** NMNS PM 65051-2, right valve; **G,** NMNS PM 65051-3, right valve. **E.** Fossil from the uppermost lower Pleistocene (Calabrian) Omma Formation (upper part) at Loc. 3, NMNS PM 65053, right valve. **H.** Fossil from the lowermost lower Pleistocene (Gelasian) Nadachi Formation (uppermost part) at Loc. 2, NMNS PM 65052, left valve. Scale bars = 10 mm.

- 1977: pl. 10, fig. 16, pl. 15, fig. 4; Fujiyama, 1982: 338, pl. 189, fig. 1633; Masuda *et al.*, 1983: 9, pl. 1, figs 11a, b, 12; Matsuura, 1985: pl. 32, fig. 17; Amano *et al.*, 1988: 68, pl. 1, fig. 7; Ozawa *et al.*, 1998: 89, pl. 16, figs 6a, b; Huber, 2010: 137; Lutaenko & Noseworthy, 2012: 29, pl. 4, figs A–H, text-fig. 10; Lutaenko *et al.*, 2019: 174, pl. 4, figs A, B.
- Anadara (Scapharca) broughtoni* — Taki & Oyama, 1954: 32, pl. 18, fig. 9, pl. 35, fig. 9; Ozaki *et al.*, 1957: 167, pl. 29, fig. 18; Ozaki, 1958: 112, pl. 21, fig. 8; Ohara, 1968: 56, pl. 2, fig. 8; Ando, 1971: pl. Q-11, fig. 2; Oyama, 1973: 79, pl. 26, fig. 8; Kanno *et al.*, 1978, pl. 1, fig. 8.
- Scapharca broughtoni* — Habe, 1965: 75, pl. 1, figs 6, 11; Habe & Ito, 1965: 111, pl. 36, fig. 1; Ishii, 1987: 6, pl. 2, figs 1a, b.
- Scapharca broughtonii* — Habe & Kosuge, 1967: 125, pl. 46, fig. 6; Kuroda *et al.*, 1971: 334, pl. 68, figs 2, 3; Habe, 1977: 36, pl. 8, fig. 2; Akamatsu, 1980: 9, pl. 4, figs 2, 4, 7, pl. 5, fig. 2; Zhao *et al.*, 1982: 88, pl. 11, fig. 8; Akamatsu, 1987: pl. 1, figs 7, 8; Qi *et al.*, 1989: 157, pl. 10, fig. 3; Okutani *et al.*, 1989: 34; Akamatsu & Suzuki, 1990: pl. 4, fig. 7; Nemoto & Akimoto, 1990: pl. 8, fig. 13; Akamatsu & Suzuki, 1992: pl. 4, fig. 6; Akamatsu, 1992: pl. 1, fig. 1; Tsuchida & Kurozumi, 1993: 7, pl. 2, fig. 4; Nakagawa *et al.*, 1993: 31, pl. 11, figs 2a, b; Matsukuma & Okutani, 2000: 853, pl. 424, fig. 36; Min, 2004: 387, figs 1241-1, 2; Nobuhara, 2003, pl. 1, fig. 10; Li, 2004: 218, pl. 115, fig. K; Xu & Zhang, 2008: 35, fig. 74; Toba, 2009, 67, fig. 6; Jimbo *et al.*, 2017: figs 3Da, b; Kawase *et al.*, 2015: pl. 21, fig. B17; Matsukuma & Okutani, 2017: 1168, pl. 467, fig. 15.
- ? *Anadara (Scapharca) satowi* (Dunker) — Omori, 1971: Q-15, fig. 6.
- Anadara broughtoni* — Scarlato, 1981: 251, fig. 166; Evseev, 1981: 122, pl. 3, fig. 3; Baba, 1990: 237, pl. 23, fig. 8; Lutaenko, 1993: 27, pl. 1, fig. 5.
- Anadara broughtonii* — Zhang *et al.*, 2016: 243, fig. 303; Lutaenko & Noseworthy, 2019: fig. 2D.
- non Scapharca broughtonii* — Okumura, 1994: 46, pl. 8, fig. 20 [= *Scapharca akitaensis* (Noda, 1966)?]

**Material examined:** Three specimens from Loc. 1.

**Description:** Shell small for species, attaining more than 60.0 mm in length, rather thin for genus, subquadrate ( $H/L = 0.78$  and  $0.85$ ), moderately inflated and inequilateral. Antero-dorsal margin straight, forming right angle with broadly arched anterior margin; postero-dorsal margin straight, forming blunt angle with subtruncate posterior margin; ventral margin broadly arcuate. Blunt ridge running from beak to postero-ventral corner. Umbo swollen and protruding above dorsal margin; beak located at anterior one-third of shell length ( $AL/L = 0.30$  and  $0.33$ ); beak prominent prosocline. Surface sculptured by 42 and 44 narrowly elevated, flat-topped radial ribs separated by nearly equal interspaces with many fine growth lines; some nodes recognized at crossing points of radial ribs and growth lines in left valve. Ligamental area rather wide, crescent shaped and sculptured with three distinct grooves and some weak horizontal grooves. Hinge line straight, with 52 and 54 small vertical teeth to hinge base. Pallial line rather deep. Inner ventral margin crenulate. Anterior adductor muscle ovate; posterior adductor muscle scar subquadrate and larger than anterior one.

**Remarks:** The Tentokuji species has a smaller shell with less inflated ventral margin than the Recent specimens, which attain 150 mm (Habe, 1965). However, it can be confidently identified as *Scapharca broughtonii* because it has the same or a similar number of radial ribs (42 and 44) as the type specimens and similar ratio of shell height to length ( $H/L = 0.78$  and  $0.85$ ;  $0.78$  Liang *et al.*, 2011, for Korean population). Moreover, the shell outline of *S. broughtonii* has a wide range of variation (Ando, 1971).

The Tsunozu Research Group (1983) recorded the occurrences of “Akagai” (the Japanese name of *S. broughtonii*) without any illustrations from the marine clay M1 of the upper Pliocene (Piacenzian) part (Kano *et al.*, 2001) and M4 clay of the lower Pleistocene (Gelasian) part (Kano *et al.*, 2001) of the Tsunozu Formation in Shimane Prefecture on the Japan Sea side of western



Honshu. Takao (1996) illustrated the outer molds of shells (pl. 1, figs 1, 2 in Takao, 1996) from the M1 clay of the Tsunozu at Ushiroji-cho, Gotsu City, Shimane Prefecture. These fossils have very few radial ribs (about 22 ribs counted from his figures) and smaller shells. We also examined four ill-preserved fossils from the upper Pliocene (Piacenzian) M1 clay of the Tsunozu Formation at Ushiroji-cho, stored at the Tokushima Prefectural Museum (TKPM-GFI6285) and identified them as probably *Scapharca kagoshimensis* (Tokunaga, 1906) because of their small shell size (43 mm) and 32 radial ribs.

*Scapharca akitaensis* (Noda, 1966) from the upper middle Pleistocene (Chibanian) Shibikawa Formation at Tayasawa in Akita Prefecture, northeast Honshu is similar to a juvenile shell of *S. broughtonii*. However, *S. akitaensis* has a small shell (21.7 mm in length) and larger number of radial ribs (48 in the left valve, 47 in the right valve) like the Chinese population of *S. broughtonii* from the Bohai Sea. *Scapharca nopporoensis* Akamatsu, 1980 from the upper middle Pleistocene (Chibanian) Otoebatsugawa Formation in central Hokkaido has a very similar shell and size (92 mm in length) to *S. broughtonii*. Like *S. akitaensis*, *S. nopporoensis* has more numerous radial ribs (46 in the left valve, 45 in the right valve) than *S. broughtonii*. Although Akamatsu (1980) did not compare his species with *S. akitaensis*, it is plausible that *S. nopporoensis* is a junior synonym of *S. akitaensis*. Okumura (1994) described *S. broughtonii* from the upper middle Pleistocene (Chibanian) Ninomiya Formation in Kanagawa Prefecture, central Honshu. According to his description, the species has a medium-sized right valve (65.5 mm) and 48 radial ribs as in *S. akitaensis*. It is possible that Okumura's species is *S. akitaensis* because of the number of radial ribs.

*Scapharca ommaensis* is the characteristic species of the Omma-Manganji fauna (Otuka, 1939a) which flourished in the Japan Sea side of Honshu and southwestern Hokkaido during the Pliocene to early Pleistocene. *S. ommaensis* resembles *S. broughtonii* in having narrowly elevated radial ribs separated by nearly equal interspaces. Unfortunately, there are few descriptions or measurements of this species. Our examination of the species from the Omma Formation (Loc. 5) shows that *S. ommaensis* differs from *S. broughtonii* by having more centrally situated beak (average of  $AL/L = 0.34$ ), lower shell (average of  $H/L = 0.78$ ) and fewer numerous radial ribs (35–40, average 37.2).

*Scapharca satowi* from Loc. 6 can be easily distinguished from *S. broughtonii* by having a thick shell, more centrally situated beak (average of  $AL/L = 0.37$ ) and fewer radial ribs (35–40, average, 37.4). Moreover, *S. satowi* has wider and smooth radial ribs with no nodes in the left valve.

**Measurements:** NMNS PM 65051-1 (left valve): Length (L) 56.7 mm, Height (H) 48.1 mm, Anterior length (AL) 17.2 mm, Number of ribs (NR) 42. NMNS PM 65051-2 (right valve): L 39.4 mm, H 30.9 mm, AL 13.2 mm, NR 44. NMNS PM 65051-3 (right valve): L 60.0 mm+, H 48.1, AL 25.4 mm, NR 39+.

**Distribution:** Late Pliocene (Piacenzian), Tentokuji Formation in Akita Prefecture (this study); earliest Pleistocene (Gelasian), Nadachi Formation in Niigata Prefecture (Amano *et al.*, 1988; this study); early Pleistocene (Gelasian), Nojima Formation (Imaizumi Member), Renkoji and Imuro Formations in Kanagawa Prefecture (Baba, 1990; Jimbo *et al.*, 2017); late early Pleistocene (Calabrian), Aburayama Formation in Shizuoka Prefecture (Ozawa *et al.*, 1998), Omma Formation (upper part) in Ishikawa Prefecture (this study); middle Pleistocene (Chibanian), Dateyama, Otoebetsugawa and Umaoi Formations in central Hokkaido (Fujie, 1958; Akamatsu, 1980, 1987; Akamatsu & Suzuki, 1990, 1992; Oka, 1992), Kiyokawa, Mandano and Sanuki Formations in Chiba Prefecture (Noda, 1966; Baba, 1990), Naganuma Formation in Kanagawa Prefecture (Yokoyama, 1920, 1922; Baba, 1990), Hamamatsu Formation in Shizuoka Prefecture (Nobuhara, 2003); Takamatsu Silty Sandstone of Atsumi Group (Hayasaka, 1961; Kawase *et al.*, 2015); late Pleistocene, Tsukabara Formation in Fukushima Prefecture (Noda, 1966), Narita Formation in Ibaraki Prefecture (Kanno *et al.*, 1978), Katori and Narita Formations in Chiba Prefecture (Yamakawa, 1911; Ozaki, 1958; Ohara, 1968), Tokyo Formation in Tokyo Metropolis

(Tokunaga, 1906; Yamakawa, 1911; Ozaki *et al.*, 1957; Ando, 1971), Hiradoko Lower Shell bed, Wakura-eki Shell Bed and Ushima Shell Bed in Ishikawa Prefecture (Matsuura, 1977, 1985); Holocene, Post-glacial deposits from off Primorie (44°N; Evseev, 1981), Yurakucho Formation in Tokyo Metropolis, Nanba Formation in Osaka Prefecture (Ishii, 1987), Takahama Shell Bed in Fukui Prefecture (Nakagawa, 1993); Recent, Peter the Great Bay, Busse Lake in southern Sakhalin (empty shell; Akamatsu, 1992), southern Hokkaido to Kyushu, Japan Sea, Korea, Yellow Sea and Bohai Sea (Higo *et al.*, 1999; Scarlato, 1981; Zhao *et al.*, 1982; Matsukuma & Okutani, 2000, 2017; Li, 2004; Xu & Zhang, 2008; Min, 2004; Lutaenko & Noseworthy, 2012, 2019; Zhang *et al.*, 2016).

***Scapharca aff. broughtonii* (Schrenck, 1867)**

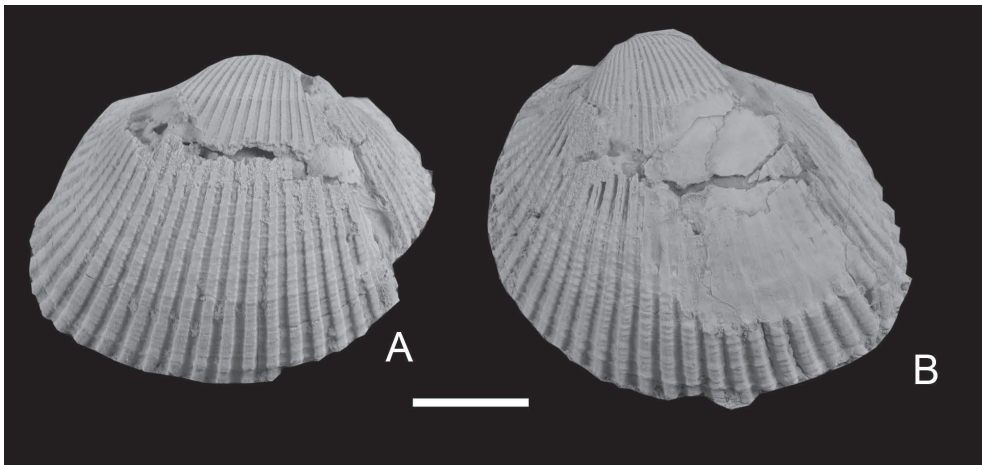
(Fig. 3)

**Material examined:** Two specimens from Loc. 4.

**Description:** Shell very small for genus, attaining 34.4 mm in length, quadrate ( $H/L = 0.78$ ,  $0.82$ ), well inflated and inequilateral. Antero-dorsal margin straight, forming blunt angle with broadly arched anterior margin; postero-dorsal margin straight, forming blunt angle with obliquely truncated posterior margin; ventral margin broadly arcuate. Blunt ridge running from beak to postero-ventral corner. Umbo well inflated and protruding above dorsal margin; beak located at anterior one-third of shell length ( $AL/L = 0.28$ ,  $0.30$ ); beak prominent prosocline. Surface sculptured by 41 and 42 narrowly elevated, flat-topped radial ribs separated by nearly equal interspaces with many fine growth lines; some nodes at crossing points with radial ribs and growth lines in anterior half of both valves. Ligament and inner structure unknown.

**Remarks:** The present species is very similar to the adult shell of *Scapharca broughtonii* in having a subquadrate outline and 41 or 42 radial ribs despite its small size, but young specimens of *S. broughtonii* have generally less inflated and elongate shells. Moreover, it is difficult to examine the inner structure. Therefore, we treat it as an uncertain species, *S. aff. broughtonii*.

**Measurements:** NMNS PM 65054-1 (right valve): Length (L) 32.5 mm, Height (H) 25.2 mm, Anterior length (AL) 9.2 mm, Number of ribs (NR) 41. NMNS PM 65054-2: (left valve): L 34.4 mm, H 28.1 mm, AL 10.4 mm, NR 42.



**Fig. 3.** *Scapharca aff. broughtonii* from the upper Pliocene (Piacenzian) Higashigawa Formation at Loc. 4. **A.** NMNS PM 65054-2, right valve. **B.** NMNS PM 65054-1, left valve. Scale bar = 10 mm.

**Distribution:** Late Pliocene (Piacenzian), Higashigawa Formation in Niigata Prefecture.

## Discussion

The Japan Sea was semi-enclosed in the early Pliocene, opening through the northern straits (Ogasawara, 1994). From the early Pliocene to early Pleistocene, the Omma-Manganji fauna mainly consisting of cold-water and endemic species flourished in the semi-enclosed sea (Ogasawara, 1986; Amano, 2001, 2007). The oldest warm-water species have been recognized from the lower Pliocene part (*ca.* 4 Ma) of the Mita Formation in Toyama Prefecture, central Honshu (Amano *et al.*, 2008). However, the Tsushima warm current reached northern Honshu around 3.5 Ma and several warm-water species invaded the cold sea (Gallagher *et al.*, 2015; Itaki, 2016; Amano, 2019a, b; Amano, 2020).

*Anadara amacula* (Yokoyama, 1925) is the best known Anadarinae of the Omma-Manganji fauna and first appeared in the late Miocene (Amano, 2001). *Scapharca ommaensis*, *S. satowi* and *S. kagoshimensis* occurred in the Japan Sea borderland beginning in the late Pliocene (Ogasawara *et al.*, 1986; Nagamori, 1998; Amano, 2001, 2007). As a result of this study, it has been revealed that *S. broughtonii* from the Tentokuji Formation in Akita Prefecture is the oldest known record of this species. The age of the formation was assessed to about 3.5 Ma, based on the calcareous nannofossils and diatom fossils (Amano *et al.*, 2000). Moreover, *S. aff. broughtonii* was recovered from the upper Pliocene Higashigawa Formation in Niigata Prefecture whose age was assessed as older than 3.3 Ma by fission track dating method (Amano, 1994). The oldest known record of *S. satowi* is from the Tentokuji Formation at Awasegai (Loc. 1 in this study) in association with *S. ommaensis* (Ogasawara *et al.*, 1986). Consequently, *S. broughtonii* and *S. satowi* first appeared in the upper Pliocene Tentokuji Formation at Awasegai of Akita Prefecture. According to the molecular phylogeny, *S. broughtonii* and *S. satowi* are more closely related than other *Scapharca* species (Matsumoto & Hayami, 2001; Feng *et al.*, 2015). The oldest known *S. ommaensis*, which was proposed as a subspecies of *S. satowi* by Otuka (1936), has been recorded from Loc. 7 of the Mita Formation (Amano *et al.*, 2008), slightly lower than MT1 Tuff (3.4–3.5 Ma by the fission-track dating method; Goto *et al.*, 2014). There is no Pliocene record of *S. broughtonii* or *S. satowi* from the Pacific side. In the late Pliocene, the Tsushima warm current reached the southwestern part of Aomori Prefecture, northernmost Honshu (Amano, 2020). In the late Pliocene, the shallow warm current affected about 50 m of the water column and sea surface temperature was 4°C higher than today (Kitamura & Kimoto, 2006; Amano, 2019a). During this time, the warm current carried the ancestral species of *S. broughtonii*, *S. satowi* and *S. ommaensis* into the originally cold-water Japan Sea where the characteristic species of the Omma-Manganji fauna lived. The genetically closely related *S. broughtonii* and *S. satowi* speciated near thenorthernmost limit of the warm current in the semi-enclosed Japan Sea. During this time, *Scapharca castellata* (Yokoyama, 1923) and *Diluvarca suzukii* (Yokoyama, 1926) dominated the warm-water Kakegawa fauna on the Pacific side of southwestern Japan (*see* Noda, 1966).

In the early Pleistocene, *S. broughtonii* occurred in the mild-temperate areas on the Japan Sea side while it was distributed in the subtropical to mild-temperate area on the Pacific side (*see* Ogasawara, 1994). The size of the shell increased in the Pleistocene. For example, the shell length of the specimen collected from the lowermost lower Pleistocene Nadachi Formation at Loc. 2 is 80.8 mm and that from the uppermost lower Pleistocene Omma Formation (upper part) at Loc. 3 is 89.0 mm.

In the middle Pleistocene, *S. broughtonii* reached the Japan Sea side of central Hokkaido. According to Fujie (1958) and Akamatsu (1984), *S. broughtonii* is associated with many cold-water species and a few temperate species such as *S. kagoshimensis* and *Phacosoma japonicum* (Reeve, 1850) from the lower part of the Dateyama Formation which corresponds to the higher terrace



deposits (Oka, 1992). Judging from the stratigraphic position and correlation with the deposits around the Ishikari Lowland, the formation can be assigned to the middle Pleistocene (Oka, 1992; Ozaki & Komatsubara, 2014). This was the earliest appearance of this species in cold-temperate water. It has been revealed that *Saxidomus purpurata* (Sowerby) adapted to the cold-temperate water in the end of the Miocene to the early Pliocene along the global climatic cooling (Amano & Nemoto, 2020). Probably there were many more cases of temperate species adapting to cold-water environments as *S. broughtonii* and *S. purpurata* did.

It is interesting to note that *Scapharca akitaensis* and *S. nopporoensis*, which resemble *S. broughtonii* but have more numerous radial ribs (45–48), occurred with warm-water species from the middle Pleistocene Otoebaetsugawa Formation in Hokkaido (around 0.4 Ma; Akamatsu & Suzuki, 1990), the Ninomiya Formation in Kanagawa Prefecture (0.451–0.265 Ma; Yano, 1986) and the Shikikawa Formation in Akita Prefecture (0.451–0.265 Ma; Sato *et al.*, 2012). In having 45 to 48 radial ribs, *S. akitaensis* and *S. nopporoensis* are similar to the Recent population of “*S. broughtonii*” in the Bohai Sea recognized by Yokogawa (1997) and Tanaka & Aranishi (2016) as a separate species or subspecies. If the Chinese population can be identified as *S. akitaensis* or *S. nopporoensis*, which appeared during the interglacial period of the middle Pleistocene in Japan, then the species would have migrated through the Japan Sea to the Bohai Sea (*see* Amano, 2005).

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## アカガイ（二枚貝綱：フネガイ科）とその近縁種の起源と生物地理学的変遷

天野和孝・小森加成江

### 要 約

アカガイ *Scapharca broughtonii* とその近縁種が本州日本海側の上部鮮新統から発見された。これはアカガイとその近縁種の最古の記録である。サトウガイ *Scapharca satowi* も秋田県の上部鮮新統天徳寺層のアカガイと同じ産地から報告されている。これらの化石記録は、遺伝的に近縁なアカガイとサトウガイが後期鮮新世までに半閉鎖的な日本海に流入した暖流の北限付近で種分化したことを示唆している。本種の冷温帯水域からの最古の記録は北海道中部の中部更新統の伊達山層からの化石である。このことは、アカガイは中期更新世に冷水域に適応進化したことを示唆している。