

# **Statistical and parametric studies on natural levees as weak points against leakages in river levees**

Wenyue Zhang<sup>a</sup> and Akihiro Takahashi<sup>a\*</sup>

*<sup>a</sup> Department of Civil and Environmental Engineering, Tokyo Institute of Technology, Tokyo, Japan*

Wenyue Zhang

Student, Department of Civil and Environmental Engineering,  
Tokyo Institute of Technology,  
ORCID: 0000-0003-2983-7376

Akihiro Takahashi\*

Professor, Department of Civil and Environmental Engineering,  
Tokyo Institute of Technology,  
2-12-1-M1-3, Oh-okayama, Meguro, Tokyo, 152-8552, Japan  
E-mail: [takahashi.a.al@m.titech.ac.jp](mailto:takahashi.a.al@m.titech.ac.jp)  
ORCID: 0000-0003-1206-5066

\* Corresponding author

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## **Statistical and parametric studies on natural levees as weak points against leakages in river levees**

Sandy natural levee deposits have been pointed out to provide seepage paths under river levees, which are largely related to the underseepage problems. However, few attempts have been taken to quantitatively study the relationship between the natural levees and the leakages in river levees. To capture the features of real natural levees, statistical studies are performed on the geometry and hydraulic conductivity of the micro-topographies along the Kinu River in Japan. By setting cases based on the retrieved data, a parametric study on the geometric and hydraulic parameters is performed by finite element seepage analysis. As a result, the embankment sitting on the landside of the natural levees is identified to be susceptible to leakages. In addition, rainfall and flooding are distinguished as the two driving forces of leakages depending on the hydraulic conductivity of the embankment bodies and the underneath foundations. The sandy natural levee deposits, with relatively high hydraulic conductivity, providing seepage paths for the under seepage, may magnify the effects of the seepage driven by the flooding, and lead to the classical backward erosion piping. Discussion and comments are addressed for the existing engineering practice in Japan.

Keywords: natural levees, micro-topographies, river levees, leakages, backward erosion piping

## 1 **1. Introduction**

2 During the 2019 Typhoon Hagibis, accompanied by which the highest daily precipitation was recorded in  
3 eastern and northern Japan, 142 cases of levee breaching occurred, causing around 64,000 hectares of  
4 submergences and 105 fatalities (FDMA, 2020; MLIT, 2020). According to the investigations, overflow,  
5 seepage flow, and scouring were recognized to be the major causes of the levee breaching (MLIT, 2019).

6 Backward erosion piping, usually referred to as “piping” or “leakage” in Japan, occurs when soil  
7 particles are detached by seepage flow at exit points, and finally leads to pipes reaching the riverside of the  
8 water retaining structures (USBR, 2019). Since the early studies by Terzaghi (1939), many studies have  
9 been performed by laboratory testing (Van Beek, 2015; Richard & Reddy, 2012; Fleshman & Rice, 2014;  
10 Negrinelli et al., 2016; Robbins et al., 2018), theoretical analyses (Sellmeijer, 1988; Rhee & Bezuijen,  
11 1992), numerical simulation (Fujisawa et al., 2010; Wang et al., 2014; Vandenboer et al., 2014; Liang et  
12 al., 2017; Maeda et al., 2019), centrifugal modelling (Van Beek et al., 2010; Ito et al., 2021), and full-scale  
13 modelling (Sellmeijer et al., 2011; PWRI, 2014; Parekh et al., 2016). Based on the knowledge from the  
14 studies, guidelines on the management of levees have been set up (USACE, 2000; USACE, 2005; JICE,  
15 2012).

16 Regretfully, with all the efforts paid, the existing evaluating methods of piping were found to be  
17 not able to correctly predict the piping failures in some cases (Kikumori, 2008). The main reason is that  
18 most of the studies were performed under idealized situations, with simplified structures and homogeneous  
19 soil. However, the real levees, constructed in different historical periods, and sitting on natural foundations,  
20 are far more complex than the idealized models. Especially, the foundations consisting of natural deposits  
21 of different micro-topographies are pointed out to play an important role in the triggering of piping (Kolb,  
22 1975; Strange et al., 2016; Dunbar et al., 2018).

23 Natural levees, as one of the commonly seen micro-topographies in the alluvial environment in  
24 Japan, are the sandy or silty structures along the river channels deposited by the overbank flow during  
25 historical floods (Brierley & Fryirs, 2005). As shown in Fig. 1, it is a common practice to build river  
26 embankments on these elevated and well-drained berms (Itsukushima, 2018). However, it is repeatedly  
27 reported (Kuroki & Shinagawa, 2018), and statistically proven (PWRI, 2010) that the embankments sitting  
28 on natural levees are closely related to leakage events. Although the sandy natural levees have long been  
29 considered to provide seepage paths for the under seepage and were emphasized in the design guideline  
30 (JICE, 2012), few attempts have been made to quantitatively study the effects.

31           The scope of this study is to quantitatively evaluate the relationship between natural levees and  
32 piping from engineering points of view. The Kinu River in Japan, where a large number of leakage events  
33 during the storm in 2015 were believed to be related to natural levees (Kuroki & Shinagawa, 2018), is  
34 selected as an example for the study (Fig. 2). To capture the features of natural levees, statistical studies  
35 are performed on the geometry and hydraulic conductivity of the micro-topographies along the Kinu River.  
36 Based on the retrieved data, a series of parametric studies on the geometric and hydraulic parameters are  
37 performed by finite element seepage analysis. Finally, discussions are made based on the results from the  
38 numerical simulations and the existing practice in Japan.

## 39 **2. Statistical studies on the geometry of the micro-topographies along the Kinu River**

40 As shown in Fig. 3, the “levee” in this study consists of the artificially built “embankment” and the naturally  
41 formed elevated “berm” underneath. This study focuses on the elevations of the structures and the two-  
42 dimensional (2D) spatial relationship between the embankments and the underneath berms.

### 43 **2.1: Source of data**

44 An online Geographical Information System (GIS) is provided by the Geospatial Information Authority of  
45 Japan (GSI) (Fig. 4), in which a Digital Elevation Model (DEM) is included. In the studied area, the DEM  
46 has 5 m-mesh and an accuracy of 0.3 m in elevation.

47           In this statistical study, the data are retrieved by the following steps:

48 (1) Selecting the positions where elevation profiles are to be retrieved in the map.

49           The positioning in this study is based on the coordination system applied for rivers in Japan (MLIT,  
50 2018), where the locations are usually presented in the form of “L/R 12.5k”, where “L” and “R” indicate  
51 left or right bank, the number is the distance measured from the estuaries, along the central line of the  
52 river, and “k” indicates kilometres.

53           In this study, data are retrieved from 0~53 km at the left and right banks along the Kinu River,  
54 with a spacing of around every 200~300 m.

55 (2) Retrieving the data by using the “sectional view” tool of the GIS.

56           This is performed from the downstream to the upstream, along the levees at the left and right banks  
57 respectively. In each focused position (data spot decided in Step 1), a path across the levee is manually  
58 selected by referring to the Elevation Map with a self-defined coloured scale and the aerial photos.  
59 Several trials may be made before the final judgement, ensuring that the selected path passes through

60 the steepest slope at the riverside and the landside, which are supposed to be the critical paths for  
61 seepage. An elevation profile across the selected path will be automatically generated, in which the  
62 focused elements can be identified, including the toes of the embankment and the edges of the berm  
63 beneath. Elevations of particular points and the widths of the focused structure can be manually read  
64 and noted down from the elevation profile.

65 (3) Integrating the information for the locations where data are retrieved.

66 The categories of micro-topographies are noted based on the Landform Classification Map for Flood  
67 Control. Notes are taken if necessary, including whether there was any leakage event during the flood  
68 in 2015 (KRDB MLIT, 2016), the existence of artificial structures, or any other abnormalities.

69 In addition to the data retrieved directly from the DEM, other data like the average elevation  
70 of the riverbed and the Highest Water Level (H.W.L.) are extracted from a report about the channel  
71 properties of the Kinu River (Research Institute of River Environment, 2009).

## 72 **2.2: The “embankment & berm” model**

73 A typical cross-section of the levees (L 23.1k) is shown in Fig. 5. The berms at the riverside are found to  
74 have stepped shapes, while the berms at the landside are found to have gentle slopes, which complies with  
75 the sedimentary mechanism of the overbank deposits and the erosional mechanism of the river.

76 According to the observations mentioned above, an idealized “embankment & berm” model is  
77 proposed, based on which the focused parameters are defined (Fig. 5). It is noticeable that the “berm” here  
78 may consist of the natural levee deposits or any other micro-topographies. In the model, the berm at the  
79 riverside is simplified to be a step, while the berm at the landside is simplified to be a slope.

80 The focused parameters in the proposed model are defined in Table 1.

## 81 **2.3: Discussion on the retrieved data**

82 By using the methodology introduced in Subsection 2.1, data are extracted from 145 spots at the left bank  
83 and 160 spots at the right bank. Among the 305 data spots, 216 of them (70.8%) are categorized as natural  
84 levees by the Landform Classification Map for Flood Control.

85 As shown in Fig. 6, the data retrieved from the DEM are not continuous. This is because in certain  
86 sections, mostly the sections where the river is directly restricted by the high terraces, no artificial  
87 embankments are found. Those sections are regarded to be out of the scope of this study.

88            Since the focus of this study is the relationship between natural levees and piping, the discussions  
89 and analysis on the spatial distribution of the natural levees and the characteristics of the river channel are  
90 not presented here. Instead, the retrieved data about the geometric characteristics of the embankments and  
91 the underneath berms will be used in Section 4 for a more realistic parametric study.

### 92 **3. Statistical study on the hydraulic conductivity of the levees along the Kinu River**

93 To have a better understanding of the hydraulic conductivity of the embankments and the natural levees, a  
94 statistical study on the levees along the Kinu River is performed.

#### 95 **3.1: Source of data**

96 Based on the *Design Guideline for Levees* (MLIT, 2017), a series of investigations were performed on the  
97 levees along the governmentally regulated rivers in Japan, during which boreholes were bored and samples  
98 were collected. Mainly based on the newly performed investigations, also involving some old documents,  
99 a database about the soil of the levees was set up, from which (1) the borehole logs, (2) the summary of soil  
100 testing results, and (3) the soil profiles are referred to in this study.

101            The borehole logs are provided in the form of ordinary borehole log sheets, in which the soil  
102 classification at a certain depth can be read. In the focused segments (0~53 km) of the Kinu River, 164  
103 borehole logs along the left bank and 180 borehole logs along the right bank are available.

104            The summary of soil testing results is provided in the form of spreadsheets, in which the basic soil  
105 properties of the samples collected at certain depths from the boreholes are provided. In the focused  
106 segments (0~53 km) of the Kinu River, 569 samples along the left bank and 571 samples along the right  
107 bank are available, while only among parts of them the hydraulic conductivities are estimated.

108            The soil profiles are provided in the form of estimated geological cross-sections as shown in Fig.  
109 7. In the soil profiles, the soil with similar properties at a certain depth is regarded as a “layer”, the naming  
110 of which indicates the formation (the capital letters, A = Alluvium, D = Diluvium, B/F = embankment/fill),  
111 the composition (the following letters, g = gravel, s = sand, c = clay, p = peat), and the sequence (the  
112 numbers) of the layer. In the focused segments (0~53 km) of the Kinu River, 33 soil profiles along the left  
113 bank and 33 soil profiles along the right bank are available, between which the spacing is around 1~2 km.

114 **3.2: Processing of the data**

115 As mentioned in Subsection 3.1, three kinds of materials are available from the database. In this statistical  
116 study, the mainly focused parameter, hydraulic conductivity is provided in the spreadsheets. However, the  
117 classification of the soil, the responding micro-topographies, and the relative positions of the layers are  
118 provided by other sources, without combining them together meaningful discussion cannot be made.  
119 Therefore, for every sample with hydraulic conductivity in the spreadsheets, the following processing is  
120 performed:

- 121 (1) Noting down the classification of the soil by referring to the borehole logs;
- 122 (2) Noting down the classification of micro-topographies in the responding spots by referring to the  
123 Landform Classification Map for Flood Control;
- 124 (3) Noting down the formation of corresponding layers (A/B/D) by referring to the soil profiles;
- 125 (4) Noting down whether the samples belong to the “seepage path” (explained below) by referring to the  
126 soil profiles.

127 Considering the conditions susceptible to leakages, the concept “seepage path” is raised here. As  
128 shown in Fig. 8, “seepage path” is defined as follows:

- 129 (1) Any continuous sandy/gravelly layers in the embankment;
- 130 (2) The continuous sandy/gravelly layers in the foundation close to the ground surface (with an  
131 impermeable blanket of less than 3 m).

132 After that, by referring to the borehole logs and the soil profiles, the parameters related to the  
133 seepage path are retrieved, including:

134 Thickness of the seepage paths  $T_S$  (m): Given that the layers are non-uniform, average values are taken  
135 among the thickness at the riverside, landside, and centre of the embankments. If there is no seepage path  
136 in the cross-section,  $T_S = 0$ .

137 Thickness of the impermeable layers above the seepage paths  $T_{NS}$  (m): The thickness at the centre of the  
138 embankment is taken. If there is no seepage path in the cross-section,  $T_{NS} = H_L$ .

139 Thickness of the covering impermeable blankets  $T_{cover}$  (m): Since the thin covering blankets are usually  
140 not recorded in the soil profiles in the database, only a few values are retrieved from the investigating report  
141 of the leakage spots observed in 2015 (KRDB MLIT, 2016).

142 **3.3: Discussion on the retrieved data**

143 In the newest guideline (JICE, 2012), in-situ tests are recommended for the foundations, and laboratory  
144 tests on the re-constituted samples are recommended for the embankment bodies. However, during the  
145 processing of the data, hydraulic conductivity in the database is found to be mainly estimated by the  
146 empirical Creager's method (728 out of 1,111 data). As a result, the discussions are mainly based on the  
147 data estimated by Creager's method, in which the estimating formulas for hydraulic conductivity  $k$  (m/s)  
148 based on the empirical data are given by Inazaki & Konishi (2010):

$$149 \quad k = \begin{cases} 0.36D_{20}^{2.368} \times \frac{1}{100}, & \text{if } D_{20} > 0.03 \\ 0.0647D_{20}^{1.885} \times \frac{1}{100}, & \text{if } D_{20} < 0.03 \end{cases} \quad (1)$$

150 where  $D_{20}$  is the 20% passing grain size (mm).

151 With the retrieved data, to illustrate the relationship between the hydraulic features of natural  
152 levees and the leakage events, the spatial distribution of hydraulic conductivity in natural levees (NL) is  
153 presented with highlights on seepage paths and leakage events. The comments and findings on the retrieved  
154 data are as follows:

155 (1) The retrieved hydraulic conductivities scatter intensely along the river, from which obvious trends  
156 cannot be identified.

157 Since the data are retrieved all along the boreholes, it is not surprising that hydraulic conductivities in  
158 the same location vary in a large range. Given the difficulty to distinguish which layers belong to the  
159 natural levee deposits, as well as the inaccuracy of the estimating methods, it is regarded to be not  
160 feasible to directly relate hydraulic conductivity and the leakages.

161 (2) Leakage events in the foundations, especially severe boiling events, tend to occur in locations with  
162 continuous seepage paths through the foundation.

163 In Fig. 9 (a), most of the leakage events in the foundations are found to be related to the seepage paths  
164 through the foundation. Therefore, the existence of continuous permeable layers through the  
165 foundation is believed to be the critical condition of leakages rather than the hydraulic conductivity  
166 itself.

167 (3) Seepage paths through the embankments are also associated with leakage events.

168 It is found in Fig. 9 (b) that the leakage events through the embankments tend to occur in locations  
169 with continuous seepage paths. However, unlike the materials in the foundations, which are believed  
170 to be related to the corresponding micro-topographies, the materials in the embankments depend on



171 the construction process, which is out of the scope of this study. Considering that the artificial  
172 embankments are built on elevated berms, another possibility is that some of the seepage paths in  
173 embankments consist of the natural materials in the foundations. It is common that the soil profiles  
174 based on the investigations after the leakages (KRDB MLIT, 2016) contradict the old documents,  
175 showing that parts of the embankment body previously regarded as artificial materials turn out to be  
176 natural materials (for example, L 20.15k and L21.5k in the Kinu River).

177 In summary, the data about the hydraulic characteristics of natural levees are retrieved, based on  
178 which the relationship between the seepage paths and the leakage events is revealed. Most importantly, the  
179 knowledge and the data achieved in the study will contribute to the parametric study in Section 4.

#### 180 **4. Parametric study on the natural levees underneath the embankments**

181 In this section, a series of parametric studies are performed to quantitatively evaluate the relationship  
182 between natural levees and piping risk. The significance of parametric studies is that the simulations are  
183 performed based on the information achieved from the statistical studies, which ensures the representation  
184 of reality.

##### 185 ***4.1: The simulated model***

186 The “Kanto-Tohoku Heavy Rainfall” in 2015 (KRDB MLIT, 2016; Technical Committee on the Kinu River  
187 Levees, 2016), which led to a series of leakage events, is taken as the prototype of the simulations in the  
188 study. Based on the knowledge from the statistical studies and the prototype, the simulated model is built,  
189 as shown in Fig. 10.

190 In this model, the geometry of the embankment is set by referring to the cross-section at L 21.0k  
191 in the Kinu River (Technical Committee on the Kinu River Levees, 2016), at which breaching occurred in  
192 the 2015 Kanto-Tohoku Heavy Rainfall, and the suggested design values by the guideline (JICE, 2012).

193 The soil profile in the model is set based on the concept of the “seepage path”. The model consists  
194 of (1) the silty embankment material Bc, (2) the sandy natural levee material As (the seepage path), (3) the  
195 clayey alluvial deposits Ac, and (4) the silty covering blanket on the surface of the embankment and the  
196 landside T. The hydraulic and mechanical parameters of the soil are set based on the field investigation at  
197 L 21.0k in the Kinu River (Technical Committee on the Kinu River Levees, 2016), as summarized in Table  
198 2. To account for the anisotropy of hydraulic conductivity, following the common practice in Japan,  
199  $k_v/k_h = 1/3$  is applied to all the soil, where  $k_v$  is the vertical hydraulic conductivity (m/s), while  $k_h$  is

200 the horizontal hydraulic conductivity (m/s) (Tanaka et al., 2017). To describe the unsaturated behaviour,  
201 the unsaturated soil property defined in the Japanese guideline (JICE, 2012; Fig. 11) is applied.

202 In the model, the focused geometric parameters (highlighted in Fig. 10) include: (1) the relative  
203 elevation of the embankment at the riverside  $H_r$ , (2) the relative elevation of the embankment at the landside  
204  $H_l$ , (3) the thickness of the seepage path  $T_s$ , and (4) the height of the seepage path above the landside  $H_s$ .  
205 The last parameter  $H_s$  (m), defined as the elevation difference between the top of the seepage path and the  
206 embankment toe at the landside, can be calculated by  $H_s = H_l - T_{NS}$ .  $H_s$  is the parameter to describe the  
207 position of the seepage path.

208 Keeping the same with the practice in Japan, the simulations in the study consider both the effects  
209 of the precipitation and the river water level rising. The applied hydraulic loading in the simulations is set  
210 based on the “Kanto-Tohoku Heavy Rainfall” in 2015 (Technical Committee on the Kinu River Levees,  
211 2016) and the Japanese guideline (JICE, 2012). As highlighted in Fig. 10, the varying head boundary is  
212 applied at the riverside; the constant head boundary is fixed at the landside; the bottom boundary is set  
213 impermeable, and the precipitation boundary is set for the unsubmerged surface. The temporal variation of  
214 the hydraulic loading is shown in Fig. 12. The hydraulic loading is divided into six stages as summarized  
215 in Table 3.

216 It is noticeable that: (1) the H.W.L. is taken as datum here so that the same hydraulic loading input  
217 can be applied for all the cases with variant geometry settings; (2) to ensure conservative situations and  
218 also to simplify the model, the river water level rises from and recovers to the elevation of the groundwater  
219 level (-7.00 m in Fig. 12), which is higher than the actual river water before and after the flooding (-10.00  
220 m and -7.40 m in Fig. 12); (3) the groundwater level is taken as the level of the lowland behind the levee,  
221 which is determined from the statistical studies as introduced in Section 2.

#### 222 **4.2: Details about the simulations**

223 In the parametric study, the finite element analysis is applied to numerically simulate the seepage process.  
224 The finite element analysis software PLAXIS 2D is controlled by the scripts written in Python so that the  
225 same template can be conveniently applied with different input geometric and hydraulic parameters  
226 (Bentley System, 2019). Transient flow analysis is performed in the plane strain model.

227 Given that the conditions with covering blankets are the more common cases in the Kinu River  
228 based on the investigations (KRDB MLIT, 2016), the index  $G/W$  is chosen as the index to describe the risk  
229 of piping, which is defined as:

230 
$$G/W = (\rho_t g \cdot H)/(\rho_w g \cdot P) \quad (2)$$

231 Where,  $G$  = weight of the covering soil,  $W$  = uplifting pressure under the covering soil,  $g$  = acceleration of  
232 gravity ( $\text{m/s}^2$ ),  $\rho_t$  = bulk density of the covering soil ( $\text{kg/m}^3$ ),  $H$  = thickness of the covering soil (m),  $\rho_w$  =  
233 density of water = 1000 ( $\text{kg/m}^3$ ), and  $P$  = pressure head under the covering soil (m).

234 As shown in Fig. 13,  $G/W$  is a kind of factor of safety against the uplifting seepage forces  
235 underneath the covering blanket. The smaller the value of  $G/W$  is, the larger the risk of piping is. A minimal  
236  $G/W$  is screened out by Python script in time and space during a simulation. A minimal value of 1.0 is  
237 required for  $G/W$  by the Japanese guideline (JICE, 2012).

### 238 *4.3: Effects of the geometry*

239 In this subsection, studies on the geometry of the embankment and the underneath seepage path are  
240 presented. Given the difficulty to study the focused geometric parameters (Fig. 10) altogether, a series of  
241 studies are conducted on the parameters separately, following the philosophy of parsimony (starting from  
242 the simplicity and building up complexity gradually). Three groups of studies are performed, focusing on  
243 (1) the effects of the thickness of the seepage paths, (2) the effects of the position of the seepage paths, and  
244 (3) the effects of the elevation difference between the landside and the riverside. It should be noted that in  
245 all the simulated cases, the setting of parameters is based on the retrieved data from the statistical studies,  
246 which are illustrated in detail in Appendix A.

#### 247 *4.3.1: Effects of the thickness of the seepage path*

248 Firstly, the focus is cast on the effects of the thickness of the seepage paths. As highlighted in Fig. 14, the  
249 simulated cases have different thicknesses of the seepage path  $T_s$  (Table 4), while all the other geometric  
250 parameters are kept constant. The hydraulic parameters of the soil follow the ones summarized in Table 2.

251 The results of the simulations are shown in Fig. 15. The x-axis in the figure is the thickness of the  
252 seepage path  $T_s$  normalized by the relative elevation of the embankment at the landside  $H_l$ . It is found that  
253  $G/W$  drops with the increasing thickness of the seepage path, indicating a larger piping risk for a thicker  
254 seepage path.

255 4.3.2: *Effects of the position of the seepage path*

256 For the embankments built on the elevated berms, it is possible that the parts that appeared to be the  
257 embankment bodies in geometry consist of the natural levee deposits. Therefore, the seepage paths may not  
258 be only located in the foundation but may also pass through the embankment bodies.

259 The study is performed on the effects of the position of the seepage path. As shown in Table 5,  
260 two groups of cases with a thickness of the seepage path  $T_S$  of 3 m and 4 m are set. In all the cases of this  
261 study, the relative elevations of the embankment at the landside and the riverside  $H_l$  and  $H_r$  are kept  
262 constant, while the positions of the seepage path are determined by the height of the seepage path above  
263 the landside  $H_S$  (Fig. 16). In all the cases, the hydraulic parameters of the soil follow the ones summarized  
264 in Table 2.

265 The simulated results are presented in Fig. 17. The x-axis in the figures is the heights of the seepage  
266 path above the ground  $H_S$  normalized by the thickness of the seepage path  $T_S$ . It is found that, although the  
267  $G/W$  values fluctuate with  $H_S$ , the variation is not considered to be large enough, especially compared with  
268 the difference caused by the thickness of the seepage path. Therefore, the position of the seepage path is  
269 not regarded as the determinant factor for the piping risk.

270 4.3.3: *Effects of the elevation difference between the landside and the riverside*

271 In the statistical study by Kuroki & Shinagawa (2018), it was concluded that leakages tend to occur at  
272 locations with riverbanks higher than the protected side behind the levees. In the model, this is  
273 corresponding to the larger relative elevation of the embankment at the landside than the elevation  
274 difference at the riverside ( $H_l > H_r$ , or  $\Delta H = H_r - H_l < 0$ ).

275 The study is performed on the elevation difference between the landside and the riverside. As  
276 shown in Table 6, two groups of cases with a relative elevation of the embankment at the landside  $H_l$  of 4  
277 m and 5 m are included. The relative elevation of the toe of the embankment at the landside to that at the  
278 riverside  $\Delta H$  is calculated for each case ( $\Delta H = H_r - H_l$ ). As illustrated in Fig. 18, the other parameters are  
279 kept constant in all the cases of this study ( $T_S = 3$  m,  $H_S = 1$  m), while the varying parameters are highlighted.  
280 In all the cases, the hydraulic parameters of the soil follow the ones summarized in Table 2.

281 The simulated results are presented in Fig. 19. The x-axis in the figures is the relative elevation of  
282 the toe of the embankment at the landside to that at the riverside  $\Delta H$  normalized by the relative elevation  
283 of the embankment at the landside  $H_l$ . It is found that:

- 284 (1) The piping risk rises drastically as the relative elevation of the embankment at the landside  $H_l$   
 285 increases. Considering that larger  $H_l$  means larger head difference, which is directly related to the  
 286 piping risk, the finding is regarded to consist with the expectation.
- 287 (2) The piping risk drops as the relative elevation of the toe of the embankment at the landside to that at  
 288 the riverside  $\Delta H$  increases. In another word, the piping risk is larger for the cases with higher  
 289 foundations at the riverside. However, the effects of  $\Delta H$  are not so significant compared to the effects  
 290 of the relative elevation of the embankment at the landside  $H_l$ , especially for the cases with relatively  
 291 large  $H_l$  values ( $H_l = 5$  m).

#### 292 4.3.4: Summary

293 In the study on the effects of the geometry of the “embankment & berm” model, it is found that:

- 294 (1) Among the focused geometric parameters, the relative elevation of the embankment at the landside  
 295  $H_l$  and the thickness of the seepage path  $T_s$  are the determinant factors of the piping risk. The piping  
 296 risk increase with the increasing  $H_l$  and  $T_s$ . The levee typically susceptible to leakages is illustrated  
 297 in Fig. 20, where the embankment is built on the landside of a berm consisting of sandy natural levee  
 298 deposits, leading to relatively large  $H_l$  and  $T_s$ . This model is similar to the cross-sections where  
 299 leakage events with severe boiling were observed along the Kinu River (KRDB MLIT, 2016).
- 300 (2) Other geometric parameters like the height of the seepage path above the ground  $H_s$  and the relative  
 301 elevation of the toe of the embankment at the landside to that at the riverside  $\Delta H$  are found not to  
 302 have determinant effects on the piping risk  $G/W$ .

### 303 4.4: Effects of the hydraulic conductivity

#### 304 4.4.1: Numerical simulations

305 In this subsection, the study is performed on the effects of the seepage paths. All the geometric parameters  
 306 are kept the same with Case 2-2 ( $T_s = 3$  m,  $H_s = 1$  m,  $H_l = H_r = 5$  m,  $\Delta H = 0$  m, as shown in Fig. 21), while  
 307 different hydraulic conductivities are applied in different cases, as summarized in Table 7. The relative  
 308 hydraulic conductivity  $R_k$  is defined as the ratio of the vertical hydraulic conductivity in the seepage path  
 309 (the As layer) to the vertical hydraulic conductivity in the embankment body (the Bc layer):

$$310 R_k = (k_v [As]) / (k_v [Bc]) \quad (3)$$

311 In Table 7, three groups of cases are set: (1) the cases with  $R_k = 1$ , (2) the cases with  $R_k = 20$ , and  
312 (3) the cases with  $R_k = 100$ . In each group, cases with different vertical hydraulic conductivity in the  
313 embankment body  $k_v$  [Bc] are included to consider the large variation in the estimated hydraulic  
314 conductivity. It is noted that the hydraulic conductivity of the T layer and the Ac layer, as well as the  
315 unsaturated characteristics of all the soil, still follow the settings in Table 2.

316 To distinguish the effects of the flooding and the rainfall, besides the simulations following the  
317 process illustrated in Subsections 4.1 & 4.2, another set of simulations are performed on all the cases,  
318 keeping all the other settings the same, but with only the rainfall as the hydraulic loading. In these  
319 simulations, the water level is kept as the initial value along the time, while the rainfall in Fig. 12 is applied.  
320 Finally, the results (index for piping risk  $G/W$ ) from the two sets of simulations (with flooding + rainfall,  
321 and with rainfall only) are compared. The small difference between the  $G/W$  values from the two sets of  
322 simulations indicates that the flooding does not contribute much to the piping risk, while the large difference  
323 indicates that the flooding is the main driving force of the piping risk.

324 In Fig. 22, the index for piping risk  $G/W$  is plotted against the vertical hydraulic conductivity in  
325 the embankment bodies  $k_v$ [Bc]. The results from the simulations with rainfall and flooding (Rain + Flood)  
326 are plotted in solid lines, while the results from the simulations with only the rainfall as hydraulic loading  
327 (Rain only) are plotted in dashed lines. It is found that:

328 (1) The distinction between the effects of flooding and rainfall can be made.

329 In the figure, the contribution of the flooding to the piping risk can be distinguished from the  
330 difference between the solid lines and the dashed lines. Looking at the lines with the same colour (the  
331 cases with the same relative hydraulic conductivity  $R_k$ ), it is found that when the hydraulic  
332 conductivity is lower than a certain level (around  $1.0E-5$  m/s for  $R_k = 1$ , around  $1.0E-6$  m/s for  $R_k =$   
333  $20$ , and around  $2.0E-7$  m/s for  $R_k = 100$ ), the solid line and the dashed line converge together, while  
334 with higher hydraulic conductivity, the solid line gradually separates from the dotted line and gives  
335 lower  $G/W$  values (larger piping risk).

336 (2) Depending on the relative hydraulic conductivity  $R_k$ , the seepage paths magnify the effects of the  
337 flooding.

338 Comparing the lines with different colours, it is found that the separating spots of the solid lines and  
339 the dashed lines differ among the groups with different  $R_k$  values. In the group with  $R_k = 1$  (in other  
340 words, no seepage path), a  $k_v$ [Bc] value of around  $1.0E-05$  m/s is needed for the seepage from the

341 riverside to contribute to the piping risk, while in the group with  $R_k = 100$  (the cases with highly  
342 seepage paths), the seepage from the riverside starts to contribute to the piping risk with  $k_v$ [Bc] value  
343 of around  $2.0E-07$  m/s.

344 To look into the mechanism behind this, the distributions of flow velocity in different cases (with  
345 flooding + rainfall) are compared in Fig. 23. This figure reveals the followings:

346 (1) The contribution of the flooding and the rainfall to the uplifting forces at the landside can be  
347 distinguished.

348 In the cases with very low transmissivity (the measure of the ability of water to transmit horizontally),  
349 like Case 4-7 and Case 4-11, the seepage from the riverside cannot penetrate the levees (embankment  
350 bodies and the foundations) and hence does not contribute to the piping risk. The accumulation of  
351 pore water pressure is mainly caused by the surface infiltration from the rainfall at the landside. In the  
352 cases with relatively large transmissivity (Case 4-10, Case 4-14), the flow patterns in the soil are  
353 dominated by the seepage from the riverside.

354 (2) The seepage paths magnify the effects of the seepage from the riverside.

355 Comparing the cases with the same  $k_v$ [Bc] but different  $R_k$  (for example, Case 2-2 and Case 4-11), it  
356 is found that the seepage paths allow the underground seepage to go further towards the landside, and  
357 hence magnify the effects of the seepage from the riverside.

#### 358 4.4.2: Discussion and comments

359 A quantitative study on the effect of hydraulic conductivity reveals the followings:

360 (1) Depending on the hydraulic conductivity of the embankments and the foundations, the leakages at the  
361 landside may be driven by: (a) the seepage from the riverside due to the flooding, (b) the surface  
362 infiltration due to the rainfall, or (c) the combined effect of the flooding and the rainfall (Fig. 24).

363 (2) The failure mechanisms are believed to be different when the driven forces are different. When the  
364 leakage is mainly driven by the seepage due to the flooding, the classical backward erosion piping  
365 may develop. In contrast, when the leakage is mainly driven by the surface infiltration due to the  
366 rainfall, the classical backward erosion piping is unlikely to develop. According to the field  
367 observations in the Kinu River (KRDB MLIT, 2016), cracks, local collapse, or slope failures are the  
368 possible consequences of the leakages driven by the rainfall. Apart from the commonly recognized  
369 sliding failure, another kind of failure called “local failure” (Akai, 1956; Wu et al., 2017; Midgley et  
370 al., 2013) or “progressive failure” (PWRI, 2015) was identified. According to some middle-scale and

371 small-scaled modelling experiments, this kind of failure occurred at the shallow part of the  
372 embankment slope and gradually progressed upward (PWRI, 2015). Up to now, there is no well-  
373 recognized estimating theory or practical regulation about “progressive failure”.

374 (3) The natural levees, which are believed to provide seepage paths for under seepage, may magnify the  
375 effects of the seepage driven by the flooding. Therefore, the embankments sitting on the natural levees  
376 are believed to be more susceptible to backward erosion piping.

377 As discussed above, the failure mechanisms may differ in the cases dominated by different driving  
378 forces. However, the evaluating methods against piping in Japan (JICE, 2012) focused on the initiation  
379 rather than the progression of piping. Therefore, the existing engineering practice is considered to be not  
380 capable of correctly capturing the “risk”, which determines the urgency of sequential countermeasures.

381 Besides the evaluating methods, appropriate countermeasures may differ for the cases dominated  
382 by different driving forces. For example, according to in-situ monitoring and large-scaled experiments  
383 (Nakata et al., 2008; Takeshita & Torigoe, 2020), the “blocking” methods like the sheet piles may be  
384 suitable against seepage from the riverside due to flooding, while the “dissipating” methods like the  
385 drainage toes may be suitable against surface infiltration due to the rainfall.

386 Given the disadvantages mentioned above, the followings are suggested:

387 (1) For the evaluating methods against piping, not only should the absolute values of the focused indexes  
388 be checked, but the driving forces of the leakages should also be distinguished. For the cases mainly  
389 driven by flooding, the classical theories about piping can be applied, while for the cases mainly driven  
390 by rainfall, a new understanding is needed.

391 (2) More accurate estimations of the hydraulic conductivity are needed so that more appropriate  
392 countermeasures can be decided based on the correct understanding of the seepage behaviour.

## 393 **5. Conclusions**

394 In this study, an attempt was made to quantitatively study the naturally formed structures. During the study,  
395 efforts are paid to adhere to reality as much as possible. Firstly, statistical studies along the Kinu River  
396 provide a basic understanding of the characteristics of natural levees. The retrieved data from the statistical  
397 studies ensure the reasonability of the simulated model and the case settings. Discussions and comments  
398 are made by associating with the engineering practice in reality. The main conclusions include:

399 (1) The assumed “embankment & berm” model is confirmed by the realistic data. The elevated berms  
400 consisting of the naturally formed alluvial deposits are identified as important parts of the levees.



- 401 (2) In the statistical study on hydraulic conductivity, the seepage paths through the embankment bodies  
402 or in the foundations are found to be closely associated with the leakage events.
- 403 (3) Among the focused geometric parameters, the relative elevation of the embankment at the landside  
404 and the thickness of the seepage path are the determinant factors of the piping risk. The embankments  
405 built on the landside of a berm consisting of sandy natural levee deposits are believed to be susceptible  
406 to leakages.
- 407 (4) Flooding and rainfall are distinguished to be the two driving forces of leakages, under the effects of  
408 which different failure mechanisms are considered to occur.
- 409 (5) The natural levees, providing the seepage paths for under seepage, may magnify the effects of the  
410 seepage driven by the flooding. Therefore, the embankments sitting on the natural levees are believed  
411 to be more susceptible to the backward erosion piping. More attention should be paid to the  
412 management of river levees in the future.
- 413 (6) This study is an attempt to combine realistic statistics and numerical analysis. A similar methodology  
414 can be applied to study other micro-topographies related to the leakages, like the abandoned river  
415 channels, back swamps, and sand dunes.

416 **Appendix A**

417 In Subsection 4.3, it is mentioned that to ensure that the simulated model correctly reflects the  
418 characteristics of natural levees, the data retrieved in the statistical studies are referred to when setting the  
419 cases in the parametric studies. In this appendix, distributions of the retrieved geometric and hydraulic  
420 parameters in the representative section (7~30 km along the left bank of the Kinu River) are presented. The  
421 representative section is selected because those natural levees are widely distributed, while cases of  
422 seepages through the seepage paths were observed in this section.

423 The distributions of the geometric and hydraulic parameters in the representative sections are  
424 shown in Fig. 25~31. In the figures, several concepts are defined to capture the characteristics of the  
425 distributions: the “range” here is the range within which all the data are distributed; the “typical values”  
426 here is the range where the data concentrate; and the “peak” here is the peak of the distribution. These  
427 features are visually distinguished from the distributions and summarized in Table 8 (Definitions of the  
428 parameters are illustrated in Fig. 5 and Fig. 8). It is seen that some of the values are blanked due to the  
429 inapplicability of the concepts to the scattered distributions, or the lack of information. Although the method  
430 seems to be tedious, it is believed to be more reasonable than trying to describe the distributions by applying  
431 a single mathematical model.

432 Besides the data from the statistics studies, data from the spots with severe sand boils are also  
433 summarized in Table 8. Retrieved from the soil profiles in the investigating reports (KRDB MLIT, 2016),  
434 these data are more accurate, with more details, and are believed to reflect the critical conditions of the  
435 failure spots. In the following parametric study, these data are taken as important references.

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### ***Competing interests***

The authors have no relevant financial or non-financial interests to disclose.

### ***Author Contributions***

Both authors contributed to the study conception and design. Data collection and numerical analysis were performed by Wenyue Zhang. The first draft of the manuscript was written by Wenyue Zhang and both authors reviewed and edited the manuscript. Both authors read and approved the final manuscript.

### ***Data availability***

The data that support the findings of this study are openly available in:

Geospatial Information Authority of Japan (GSI) at

<https://www.gsi.go.jp/top.html>

National Institute for Land and Infrastructure Management (NILIM) at

[http://www.nilim.go.jp/lab/fbg/download/geo\\_download/geo\\_download.html](http://www.nilim.go.jp/lab/fbg/download/geo_download/geo_download.html)

## Tables

**Table 1** Definition of the geometric parameters of the “embankment & berm” model.

| Parameter      | Name   | Definition   |
|----------------|--|--|
| $H_r$ (m)      | Relative elevation of the embankment at the riverside                                    | $H_r$ is defined as the elevation difference between the top of the embankment (Point 3 in Fig. 7) and the toe of the embankment (Point 2) at the riverside.   |
| $H_l$ (m)      | Relative elevation of the embankment at the landside                                     | Similar to $H_r$ , $H_l$ is defined as the elevation difference between Points 3 & 4 at the landside.  |
| $\Delta H$ (m) | Relative elevation of the toe of the embankment at the landside to that at the riverside | $\Delta H$ is defined as the elevation difference between the toe of the embankment at the landside and the riverside (Points 4 & 2), which can also be calculated by $\Delta H = H_r - H_l$ . $\Delta H > 0$ indicates that the toe of the embankment at the landside is higher than that at the riverside. |
| $T_r$ (m)      | Berm thickness at the riverside  | The thickness is defined as the elevation difference between the toe of the embankment (Point 2) and the toe of the berm (Point 1).  |
| $T_l$ (m)      | Berm thickness at the landside   | Similar to $T_r$ , $T_l$ is defined as the elevation difference between Points 4 & 5 at the landside.  |
| $W_r$ (m)      | Berm width at the riverside  | The width is defined as the distance between the toe of the embankment and the toe of the berm at the riverside (Points 2 & 1). In the case where there is no berm underneath the embankment, or the embankment locates at the edge of the berm at the riverside, $W_r = 0$ .                                |
| $W_l$ (m)      | Berm width at the landside   | Similar to $W_r$ , $W_l$ is defined as the distance between Points 4 & 5 at the landside.  |
| $W_e$ (m)      | Embankment width   | The width of the embankment is defined as the distance between the toe of the embankment at the riverside and the landside (Points 2 & 4).   |
| $\alpha_r$ (°) | Slope of the berm at the riverside   | The parameter is defined to describe the steepness of the berm at the riverside. $\alpha_r = \arctan(T_r/W_r)$ . In the case where there is no berm underneath the embankment, or the embankment locates at the edge of the berm at the riverside ( $W_r = 0$ ), $\alpha_r = 0$ .                            |
| $\alpha_l$ (°) | Slope of the berm at the landside  | Similar to $\alpha_r$ , $\alpha_l = \arctan(T_l/W_l)$ . $\alpha_l = 0$ when $W_l = 0$ .  |

**Table 2** Hydraulic parameters of the soil in the simulated model.

| Soil layer | Saturated unit weight $\gamma_{sat}$ (kN/m <sup>3</sup> ) | Vertical hydraulic conductivity $k_v$ (m/s) | Unsaturated characteristics |
|------------|---|---|-----------------------------|
| Bc         | 18  | 2.0E-06                                     | [M], [C]                    |
| As         | 17  | 4.0E-05                                     | [SF]                        |
| Ac         | 18  | 1.0E-07                                     | [M], [C]                    |
| T (Cover)  | 18  | 2.0E-06                                     | [M], [C]                    |



**Table 3** Six stages of the hydraulic loading.

| Stage | Hydraulic loading                        | Duration   |
|-------|--|------------|
| I     | Long-term rainfall                       | 0~187 hr   |
| II    | Intense rainfall                         | 187~210 hr |
| III   | Intense rainfall +<br>Water level rising | 210~224 hr |
| IV    | Intense rainfall +<br>H.W.L.             | 224~228 hr |
| V     | Water level dropping                     | 228~245 hr |
| VI    | After the flood                          | 245~250 hr |

**Table 4** Case settings in the study on the effects of the thickness of the seepage path.

| Cases | $T_s$ (m) |
|-------|-----------|
| 1-1   | 1         |
| 1-2   | 2         |
| 1-3   | 3         |
| 1-4   | 4         |
| 1-5   | 5         |
| 1-6   | 8         |
| 1-7   | 10        |

**Table 5** Case settings in the study on the effects of the position of the seepage path.

| Cases | $T_s$ (m) | $H_s$ (m) |
|-------|-----------|-----------|
| 1-3   | 3         | 0         |
| 2-1   | 3         | 0.5       |
| 2-2   | 3         | 1         |
| 2-3   | 3         | 1.5       |
| 2-4   | 3         | 2         |
| 1-4   | 4         | 0         |
| 2-5   | 4         | 0.5       |
| 2-6   | 4         | 1         |
| 2-7   | 4         | 1.5       |
| 2-8   | 4         | 2         |
| 2-9   | 4         | 3         |

**Table 6** Case settings in the study on the effects of the elevation difference between the landside and the riverside.

| Cases | $H_r$ (m) | $H_l$ (m) | $\Delta H$ (m) |
|-------|-----------|-----------|----------------|
| 3-1   | 3         | 4         | -1             |
| 3-2   | 4         | 4         | 0              |
| 3-3   | 5         | 4         | 1              |
| 3-4   | 6         | 4         | 2              |
| 3-5   | 7         | 4         | 3              |
| 2-2   | 5         | 5         | 0              |
| 3-6   | 6         | 5         | 1              |
| 3-7   | 7         | 5         | 2              |

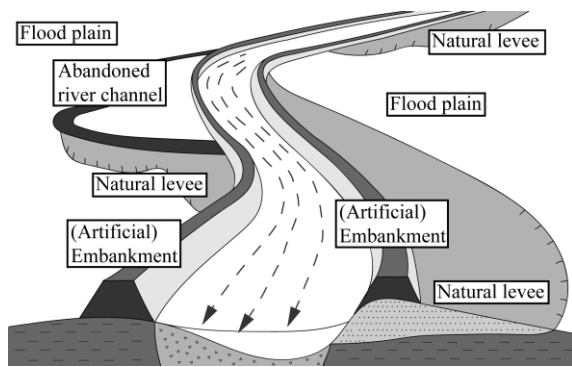
**Table 7** Case settings in the study on the effects of the seepage paths.

| Case | $R_k$ | $k_v$ [Bc] (m/s) |
|------|-------|------------------|
| 4-1  | 1     | 2.0E-07          |
| 4-2  | 1     | 1.0E-06          |
| 4-3  | 1     | 2.0E-06          |
| 4-4  | 1     | 1.0E-05          |
| 4-5  | 1     | 3.0E-05          |
| 4-6  | 1     | 1.0E-04          |
| 4-7  | 20    | 2.0E-07          |
| 4-8  | 20    | 1.0E-06          |
| 4-9  | 20    | 1.4E-06          |
| 2-2  | 20    | 2.0E-06          |
| 4-10 | 20    | 1.0E-05          |
| 4-11 | 100   | 2.0E-07          |
| 4-12 | 100   | 4.5E-07          |
| 4-13 | 100   | 1.0E-06          |
| 4-14 | 100   | 2.0E-06          |

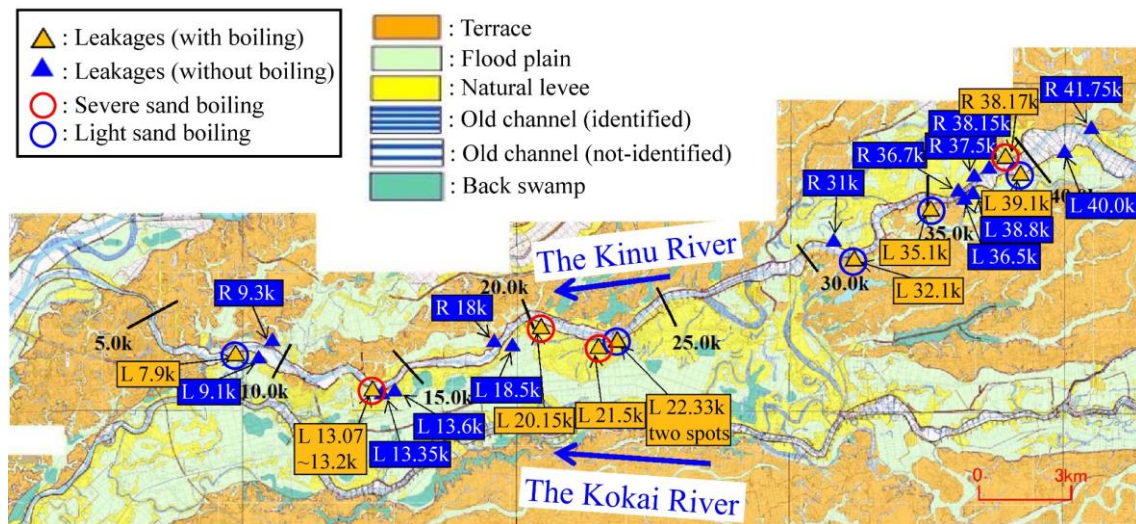
**Table 8** Features of data distributions in the representative section and the values in the spots with severe sand boils.

| Parameters      | From the statistical studies |      |                |      | From the spots with severe sand boils<br>(Technical Committee on the Kinu River Dike, 2016; KRDB MLIT, 2016) |                   |          |          |         |
|-----------------|------------------------------|------|----------------|------|--|-------------------|----------|----------|---------|
|                 | Range                        |      | Typical values |      | Peak   | Coordination (km) |          |          |         |
|                 | Min                          | Max  | Min            | Max  |  | L 13.07-13.2k     | L 20.15k | L 20.27k | L 21.5k |
| $H_r$ (m)       | 2.5                          | 9.0  | 3.0            | 7.0  | 5.0  | 4.3               | 6.8      | 5.2      | 5.5     |
| $H_l$ (m)       | 2.0                          | 7.0  | 3.2            | 5.5  | 4.2  | 3.4               | 4.8      | 3.8      | 4.4     |
| $\Delta H$ (m)  | -1.8                         | 4.0  | -0.75          | 1.0  | 0.0  | 0.9               | 2.0      | 1.5      | 1.1     |
| $\alpha_r$ (°)  | 0.0                          | 16.2 | 0.0            | 5.5  | 0.1  | 17                | 4.8      | 6.8      | 5.4     |
| $\alpha_l$ (°)  | 0.0                          | 4.0  | 0.10           | 0.60 | 0.30   | 0.17              | 0.10     | 0.16     | 0.11    |
| $T_r$ (m)       | 0.0                          | 5.2  | 2.0            | 4.0  | 2.4  | 3.0               | 0.9      | 2.4      | 1.9     |
| $T_l$ (m)       | 0                            | 4.5  | 1.8            | 3.8  | 2.6  | 3.6               | 2.0      | 2.6      | 2.8     |
| $W_r$ (m)       | 5                            | 325  | 5              | 55   | 12   | 10                | 5.0      | 20       | 20      |
| $W_l$ (m)       | 0                            | 1425 | 200            | 475  | 362  | 700               | 1170     | 900      | 1400    |
| $T_S$ (m)       | 0                            | 11.5 | --             | --   | --   | 2.7               | 2.8      | 2.3      | 2.2     |
| $T_{NS}$ (m)    | 0                            | 9.4  | --             | --   | 4.5  | 2.7               | 3.7      | 2.4      | 4.0     |
| $T_{cover}$ (m) | --                           | --   | --             | --   | --   | 0.25              | 0.20     | 0.20     | 0.70    |

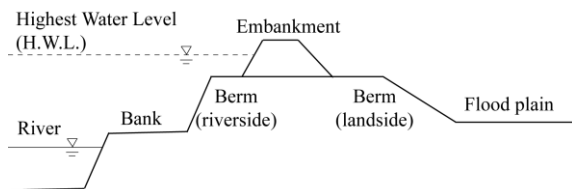
**Figures**



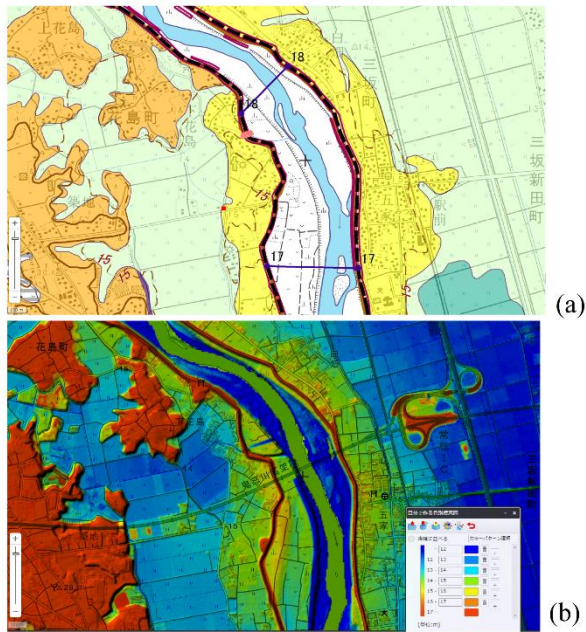
**Figure 1.** Schematic illustration of natural levees in the floodplain.



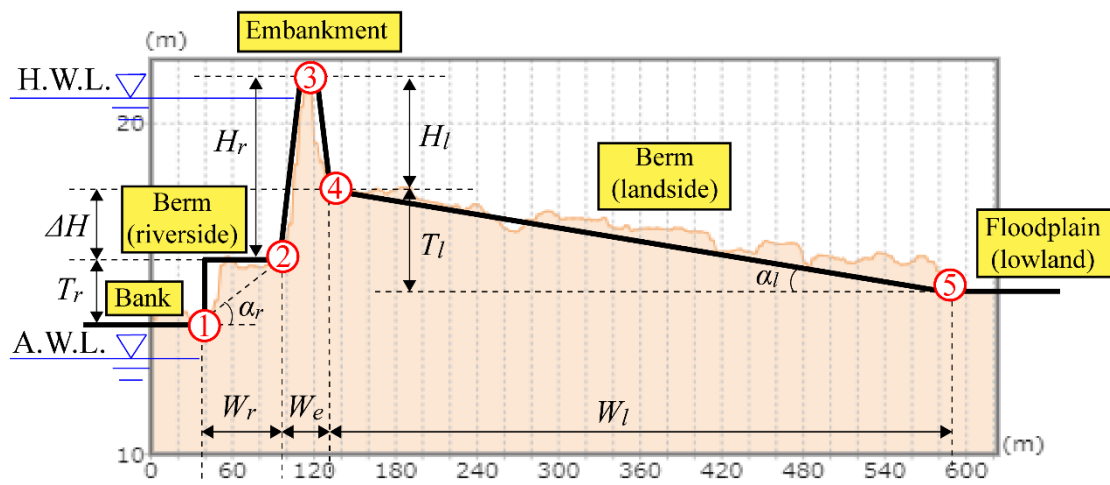
**Figure 2.** Distribution of the leakage events along the Kinu River after the “Kanto-Tohoku Heavy Rainfall” in 2015 (after KRDB MLIT, 2016).



**Figure 3.** Conceptual model in the study.

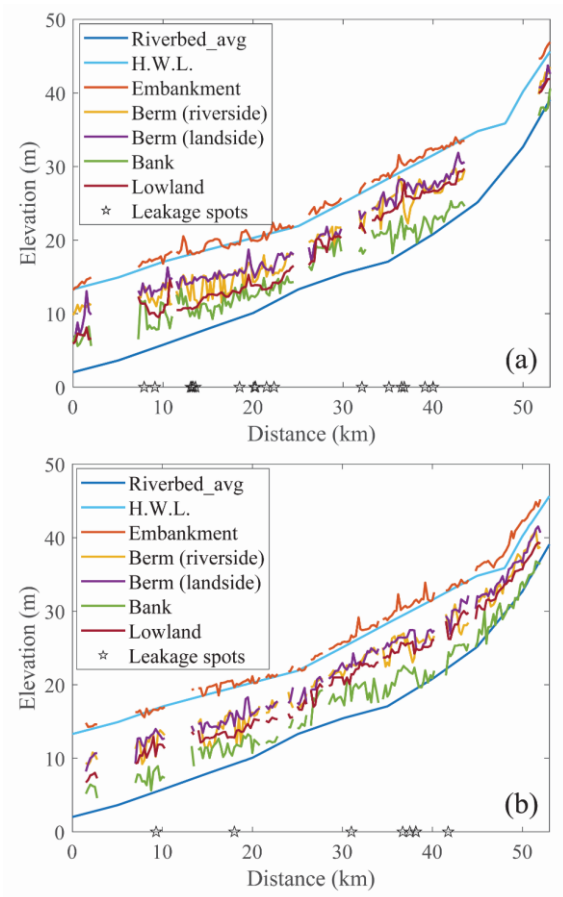


**Figure 4.** Different maps provided by the online GIS: (a) Landform Classification Map for Flood Control, (b) Elevation Map with self-defined coloured scale.

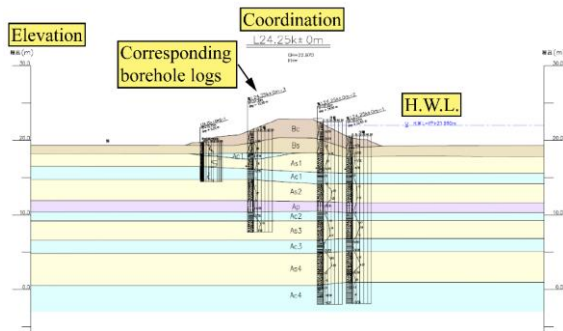


**Figure 5.** Schematic illustration of the proposed “embankment & berm” model.

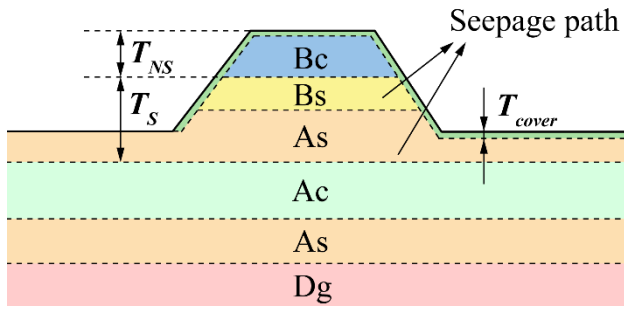




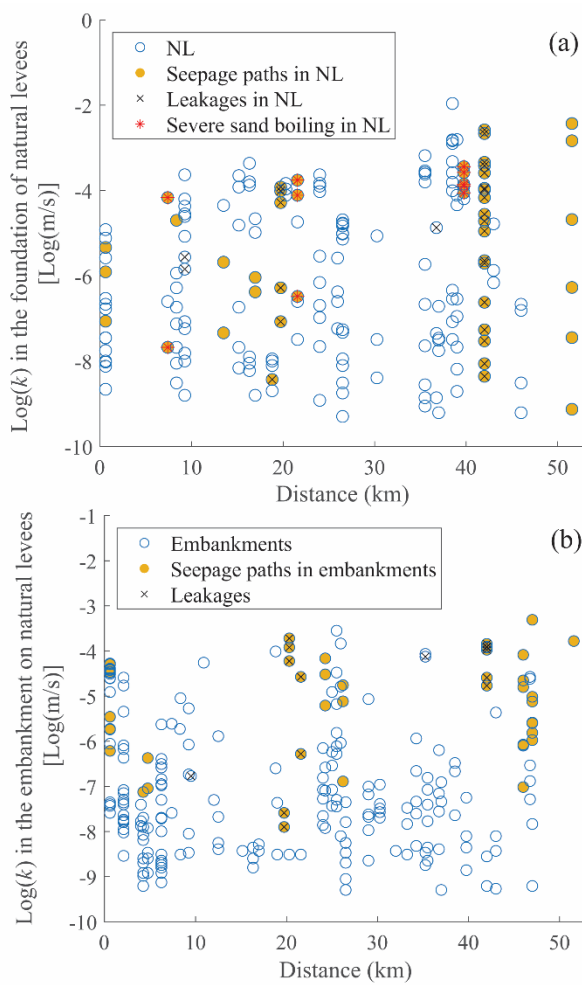
**Figure 6.** Elevation profiles along the river: (a) left bank, (b) right bank.



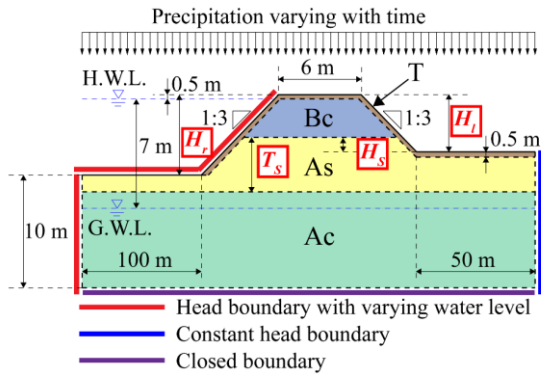
**Figure 7.** Typical soil profile in the database.



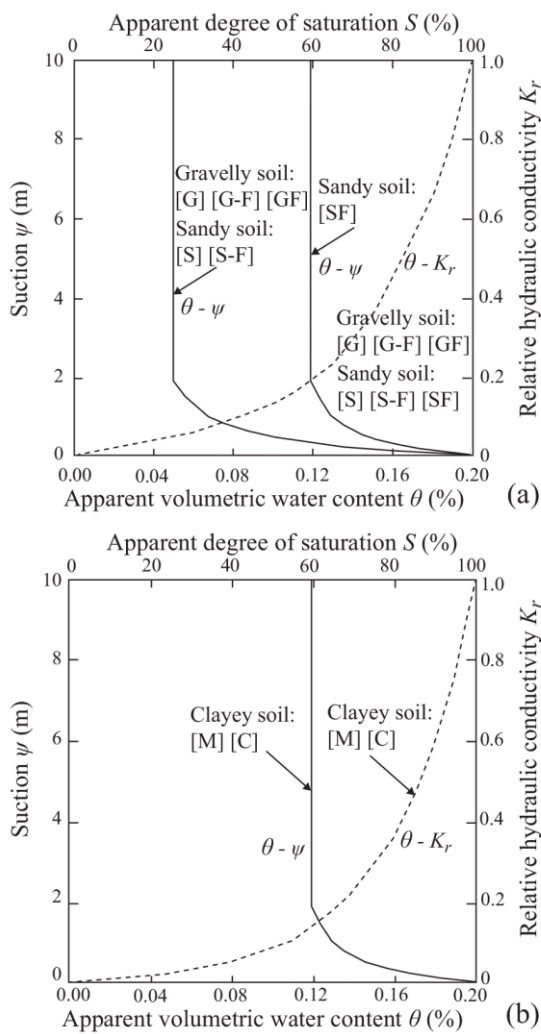
**Figure 8.** Definition of the “seepage path”.



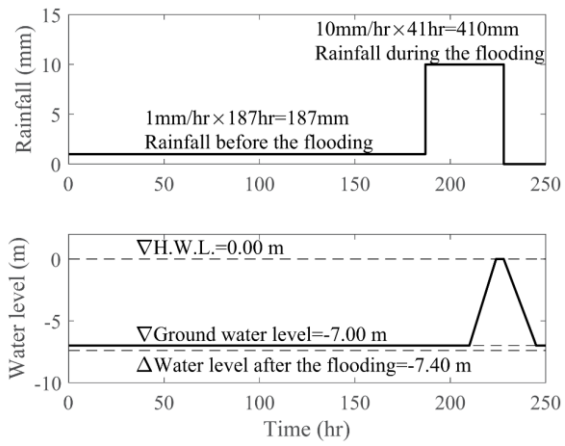
**Figure 9.** Spatial distributions of hydraulic conductivity (a) in the foundations of natural levees, and (b) in the embankments sitting on natural levees.



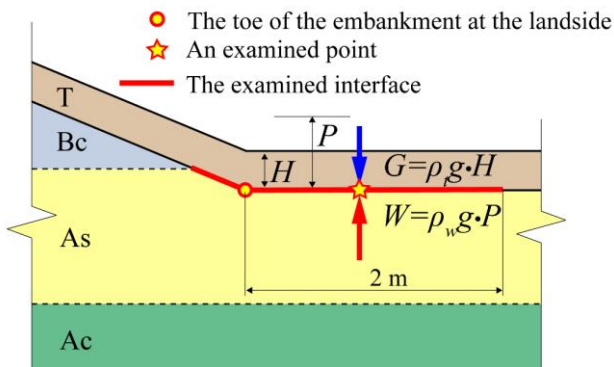
**Figure 10.** Schematic sketch of the simulated model (not in scale).



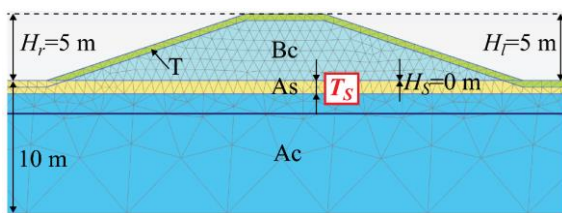
**Figure 11.** Unsaturated soil property for (a) gravelly and sandy soil, and (b) clayey soil (after JICE, 2012).



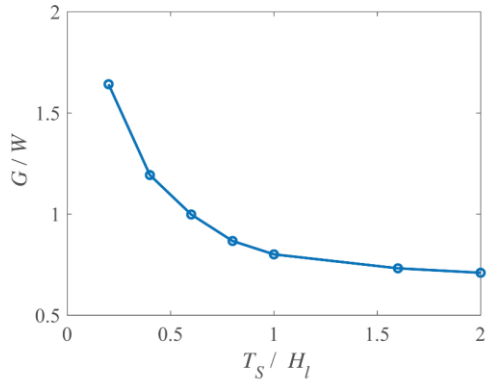
**Figure 12.** Input hydraulic loading of the simulated model.



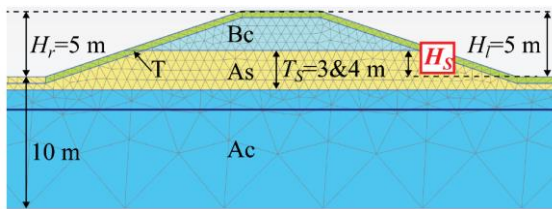
**Figure 13.** Illustration for the calculation of the index for piping  $G/W$ .



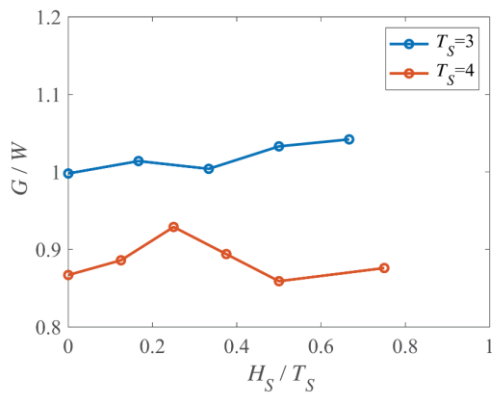
**Figure 14.** Simulated models in the study on the effects of the thickness of the seepage path (The snapshot is retrieved from Case 1-1).



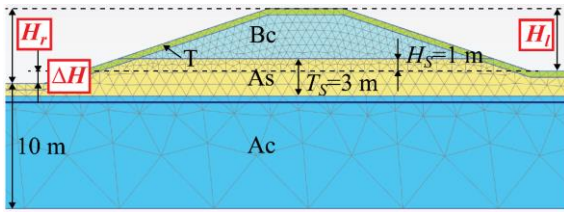
**Figure 15.** Index for piping risk  $G/W$  varying with the thickness of the seepage path  $T_S$ .



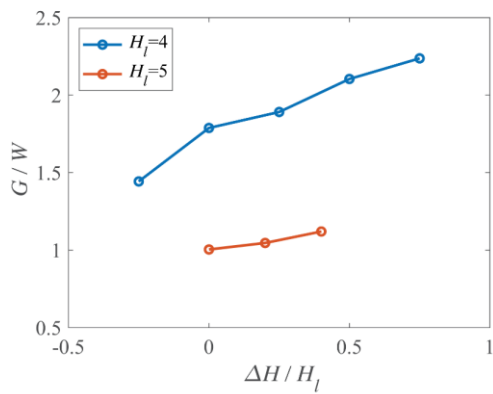
**Figure 16.** Simulated models in the study on the effects of the position of the seepage path (The snapshot is retrieved from Case 2-4).



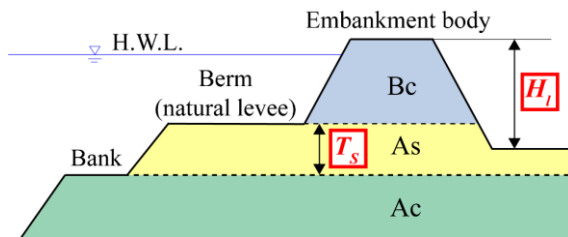
**Figure 17.** Index for piping risk  $G/W$  varying with the heights of the seepage path above the ground  $H_s$ .



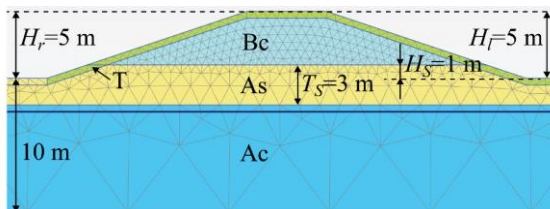
**Figure 18.** Simulated models in the study on the effects of the elevation difference between the landside and the riverside (The snapshot is retrieved from Case 3-6).



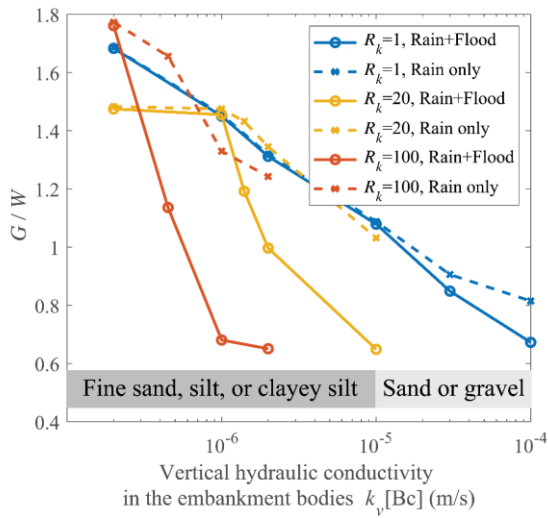
**Figure 19.** Index for piping risk  $G/W$  varying with the relative elevation of the toe of the embankment at the landside to that at the riverside  $\Delta H$ .



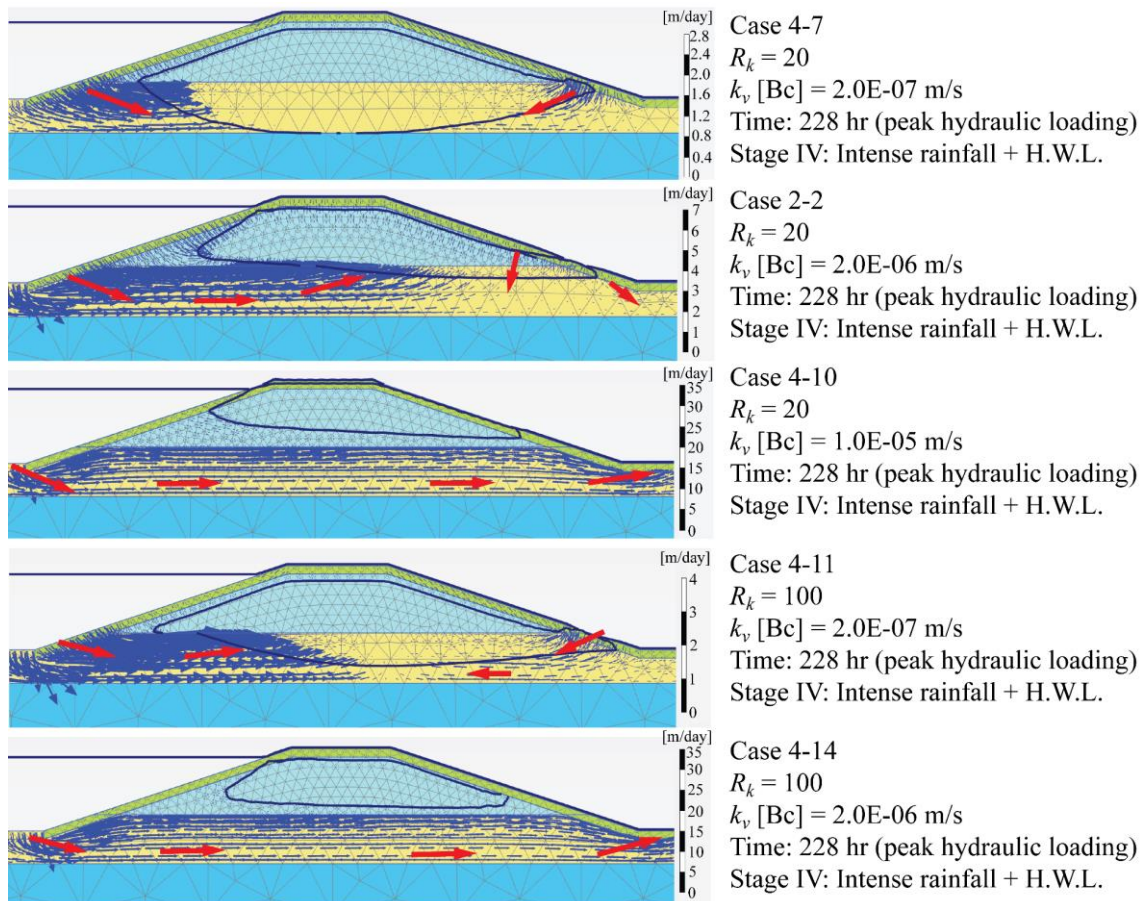
**Figure 20.** Typical geometry susceptible to leakages.



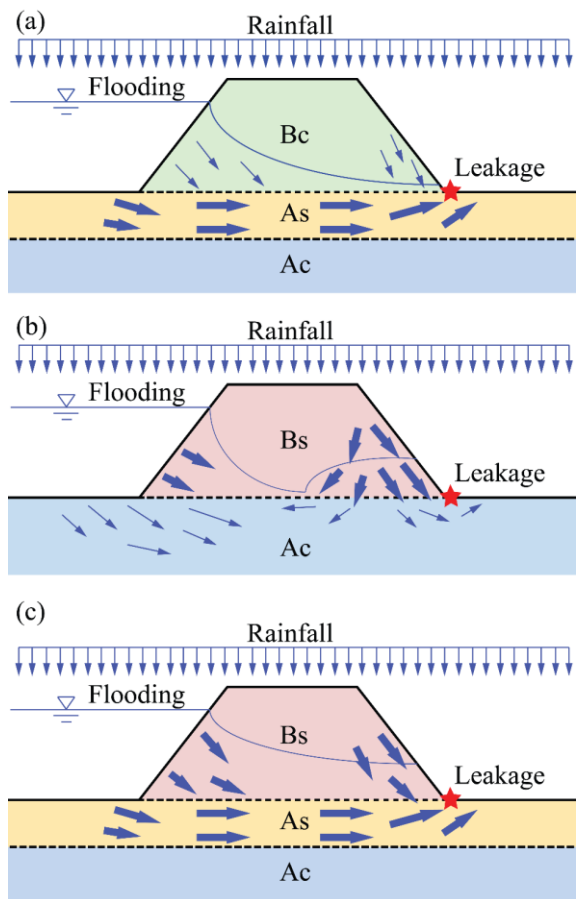
**Figure 21.** Simulated models in the study on the effects of the seepage paths (The snapshot is retrieved from Case 3-2).



**Figure 22.** Index for piping risk  $G/W$  varying with the vertical hydraulic conductivity in the embankment bodies  $k_v$ [Bc].

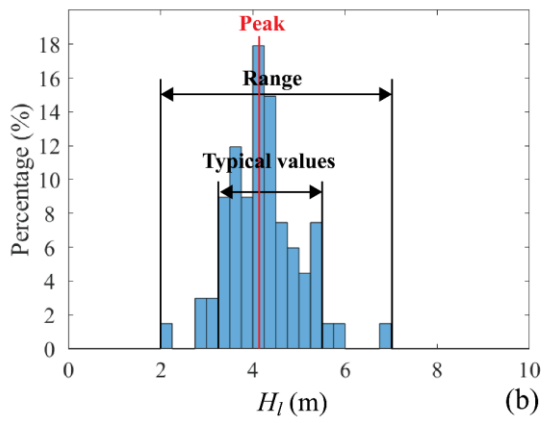
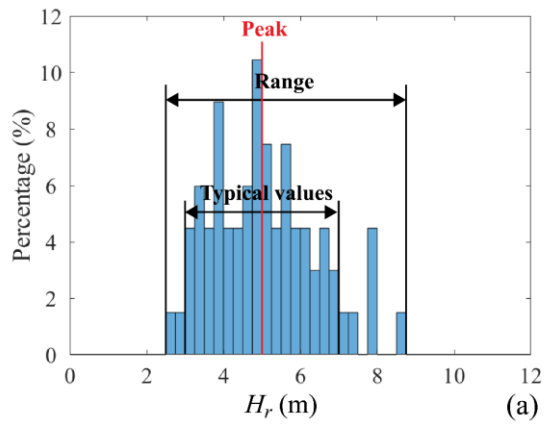


**Figure 23.** Comparison between the distributions of flow velocity under the peak hydraulic loading.

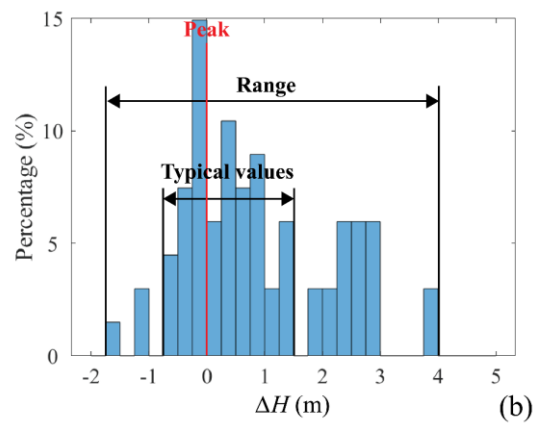
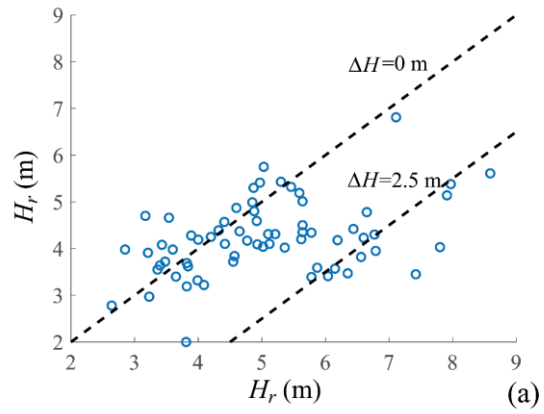


**Figure 24.** Schematic illustration of the leakage driven by (a) the seepage flow due to the flooding, (b) the surface infiltration due to the rainfall, and (c) the combined effect of the flooding and the rainfall.

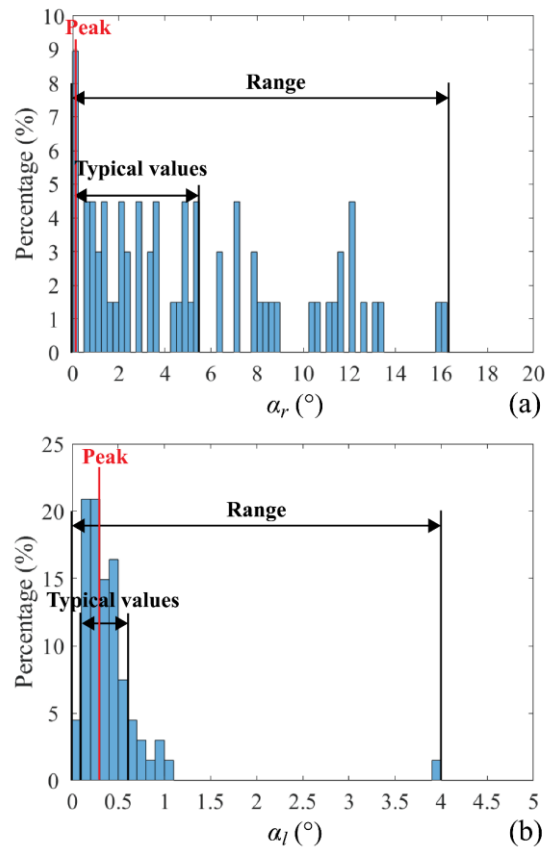




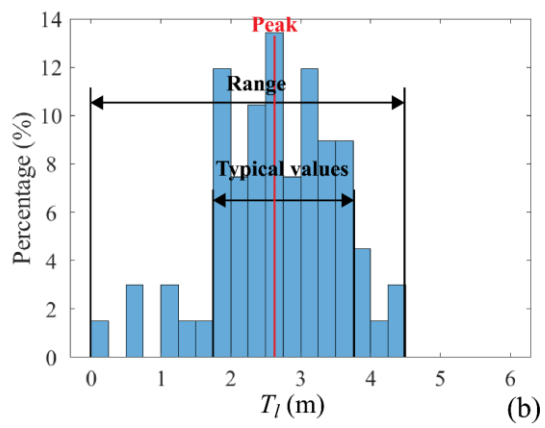
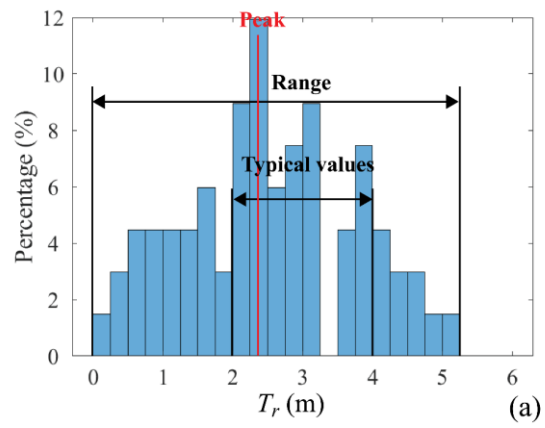
**Fig. 25** Distributions of the (a) relative elevation of the embankment at the riverside  $H_r$  and (b) the relative elevation of the embankment at the landside  $H_l$  in the representative section.



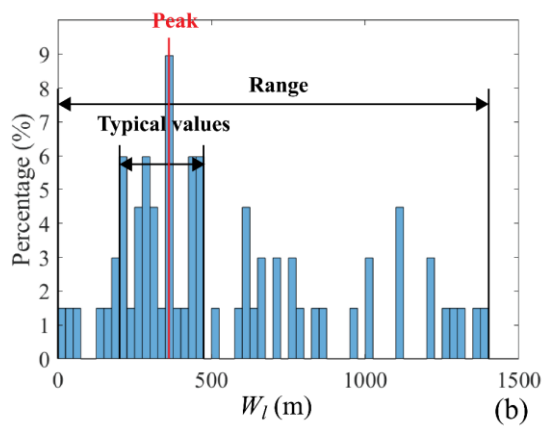
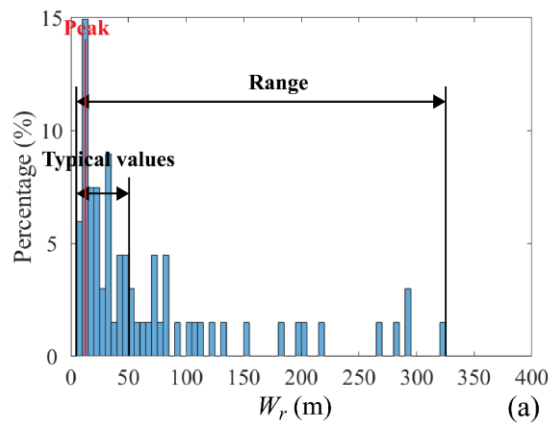
**Fig. 26** (a) Combination of the relative elevation of the embankment at the riverside  $H_r$  and that at landside  $H_l$  in the representative section; (b) Distribution of the relative elevation of the toe of the embankment at the landside to that at the riverside  $\Delta H$  in the representative section.



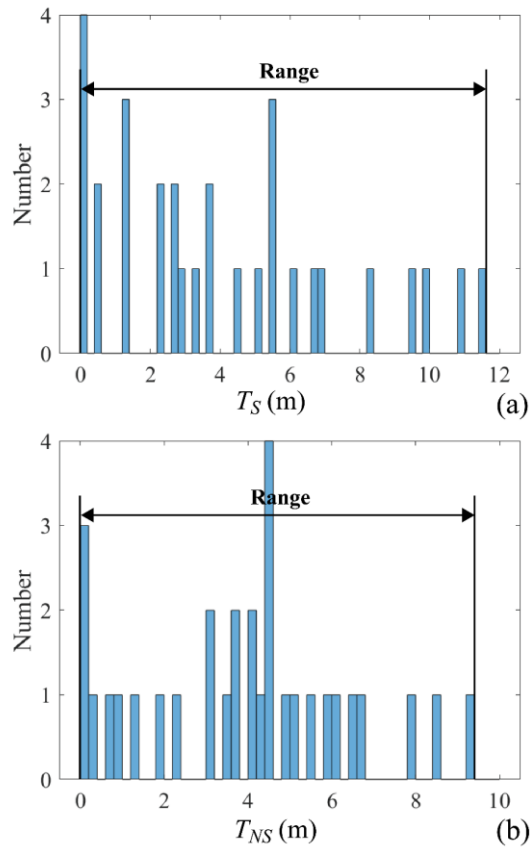
**Fig. 27** Distributions of the slope of the berm  $\alpha$  at the (a) riverside and (b) the landside in the representative section.



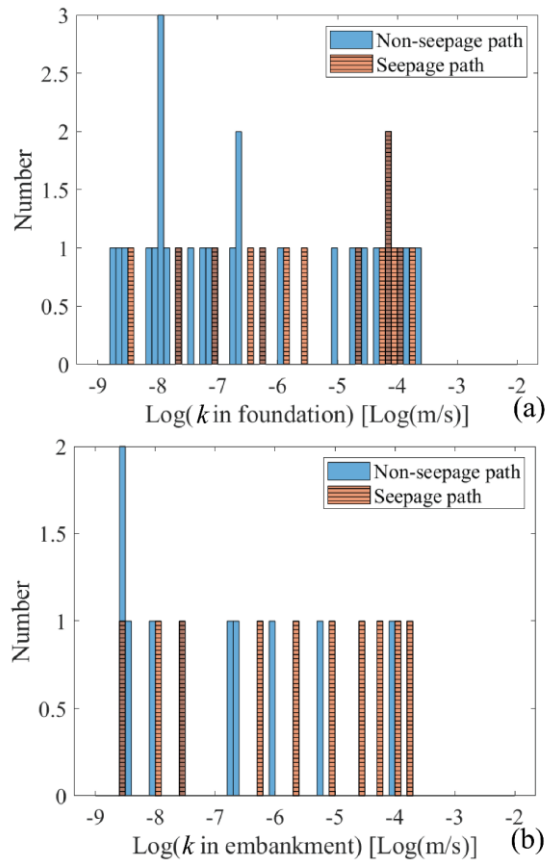
**Fig. 28** Distributions of the berm thickness  $T$  at the (a) riverside and (b) the landside in the representative section.



**Fig. 29** Distributions of the berm width  $W$  at the (a) riverside and (b) the landside in the representative section.



**Fig. 30** Distribution of the (a) thickness of the seepage paths  $T_S$  and (b) thickness of the non-seepage paths in the representative section.



**Fig. 31** Distribution of the hydraulic conductivity  $k$  estimated by Creager's method (a) in the foundations and (b) in the embankments in the representative section.